The thermal structure and the location of the snow line in the protosolar nebula: axisymmetric models with full 3-D radiative transfer

M. Min\textsuperscript{a}, C.P. Dullemond\textsuperscript{b}, M. Kama\textsuperscript{c}, C. Dominik\textsuperscript{c,d}

\textsuperscript{a}Astronomical Institute Utrecht, Utrecht University, P.O. Box 80000, NL-3508 TA Utrecht, The Netherlands
\textsuperscript{b}Max Planck Institut für Astronomie, Königstuhl 17, 69117 Heidelberg, Germany
\textsuperscript{c}Astronomical Institute 'Anton Pannekoek', Science Park 904, NL-1098 XH Amsterdam, The Netherlands
\textsuperscript{d}Department of Astrophysical IMAPP, Radboud University Nijmegen, P.O. Box 9010, 6500 GL Nijmegen, The Netherlands

Abstract

The precise location of the snow line in the protosolar nebula has been a debate for a long time. Its importance stems from the expected substantial jump in the abundance of solids beyond the snow line, which is conducive to planet formation, and from the higher stickiness in collisions of ice-coated dust grains, which may help the process of coagulation of dust and the formation of planetesimals. In an optically thin nebula, the location of the snow line is easily calculated to be around 3 AU, subject to brightness variations of the young Sun. However, in its first 5 to 10 million years, the solar nebula was optically thick, implying a smaller snowline radius due to shielding from direct sunlight, but also a larger radius because of viscous heating. Several models have attempted to treat these opposing effects. However, until recently treatments beyond an approximate 1+1D radiative transfer were unfeasible. We revisit the problem with a fully self-consistent 3D treatment in an axisymmetric disk model, including a density-dependent treatment of the dust and ice sublimation. We find that the location of the snow line is very sensitive to the opacities of the dust grains and the mass accretion rate of the disk. We show that previous approximate treatments are quite efficient at determining the location of the snow line if the energy budget is locally dominated by viscous accretion. Using this result we derive an analytic estimate of the location of the snow line that compares very well with results from this and previous studies. Using solar abundances of the elements we compute the abundance of dust and ice and find that the expected jump in solid surface density at the snow line is smaller than previously assumed. We further show that in the inner few AU the refractory species are also partly evaporated, leading to a significantly smaller solid state surface density in the regions where the rocky planets were formed.

Keywords: accretion; Solar nebula; radiative transfer; planetary formation

Email address: \texttt{M.Min@uu.nl} (M. Min)
1. Introduction

The efficiency of the formation of rocky planets and the cores of gas giant planets depends sensitively on the amount of solids present in each location of the solar nebula. In most models of planet formation the solar nebula has been assumed to resemble the structures proposed by Weidenschilling (1977) and Hayashi (1981). In this model, and subsequent incarnations of it, the surface density of the gas is approximately given by

\[ \Sigma_{\text{gas}}(r) = 1700 \left( \frac{R}{1 \text{ AU}} \right)^{-3/2} \text{gram/cm}^2 \]  

(1)

(see Thommes and Duncan, 2006) where the acronym MMSN stands for “minimum mass solar nebula”, a name by which this model is often identified. In the inner regions (small values of \( R \)) of the nebula the dust grains are refractory in nature, while beyond some radius \( R_{\text{ice}} \) water ice is present, most likely in the form of ice mantels surrounding the refractory grains. In the Hayashi MMSN model this location is assumed to be \( R_{\text{ice}} = 2.7 \text{ AU} \), which is consistent for temperatures expected in an optically thin nebula. Often this water ice sublimation boundary radius is called the ‘snow line’. Since ices may contain a considerable mass compared to the mass in refractory grains, the total surface density was assumed to jump at \( R = R_{\text{ice}} \):

\[ \Sigma_{\text{solid}}(r) = 7.1 F_{\text{ice}} \left( \frac{R}{1 \text{ AU}} \right)^{-3/2} \text{gram/cm}^2 \]  

(2)

with

\[ F_{\text{ice}} = \begin{cases} 1, & \text{for } R < R_{\text{ice}} \\ 4.2, & \text{for } R > R_{\text{ice}} \end{cases} \]  

(3)

(Thommes and Duncan, 2006)

This model serves in many planet formation models to compute the midplane gas and solid density and the isolation mass of planetary embryos. However, it has been shown that the outcomes of planet formation models depends critically on the value of \( R_{\text{ice}} \) and the strength of the ice jump \( F_{\text{ice}} \) (see e.g. Mordasini et al., 2009). Also dust coagulation and planet formation and migration models are very sensitive to these parameters (e.g. Brauer et al., 2008; Ida and Lin, 2008). Simply taking the values to be \( R_{\text{ice}} = 2.7 \text{ AU} \) and \( F_{\text{ice}} = 4.2 \) (for \( R > R_{\text{ice}} \)) is probably too simplistic and may lead to major errors in predictions of the progress of the planet formation process.

Besides the amount of solid mass available also the strength of the bonds between dust grains is of importance to their aggregation behavior. The presence of ice coatings on grains can significantly increase their ‘stickiness’, allowing grains to coagulate more easily. Perhaps even more importantly the bonds formed are stronger preventing aggregates that formed to be destroyed again by collisions.

Over the last decade or so, there has been enormous progress in astronomy in the understanding of the structure and evolution of ‘protoplanetary disks’, i.e. the dust+gas disks surrounding many very young stars. These disks very likely give a good impression of what our own solar nebula looked like 4.567 billion years ago, and the knowledge gained in that field can help getting a better handle on the structure of the solar nebula and the location of the ice condensation front. Indeed, modeling tools that were
originally meant for modeling protoplanetary disks and comparing model predictions to observations have been used to make models of the solar nebula and the snow line. For instance, the 1+1D disk structure model of Huré (2000), in which the equations are first solved vertically (1D) and then connected radially (1+1D), has been used by Hersant et al. (2001) to study water in the solar nebula, in particular concerning the issue of the D/H ratio in meteorites. Davis (2005b) made a model of the surface density of the solar nebula disk as a function of radial coordinate and accretion rate. He used this model to study the ice condensation front (Davis 2005a). This model includes both the effect of irradiation of the disk by the young sun and the heating of the disk near the midplane through viscous dissipation of potential energy (i.e. accretional heating). Like in the Hersant study, these models are 1+1D models, i.e. they are 1-D vertical models of density and temperature as a function of \( z \). The Davis models include radiative transfer using a variable eddington factor method. Lecar et al. (2006) also modeled this, using an updated version of the Chiang and Goldreich (1997) two-layer disk model, i.e. slightly simpler than the 1+1D modeling of Davis. Recently, Dodson-Robinson et al. (2009) combined complex models of ice formation and viscous evolution of the disk with a simple radiative transfer approximation to compute the evolution of the solid surface density.

While these modeling efforts have improved the understanding of the distribution of ices in the solar nebula substantially, they still rely on rather dramatic simplifications of the treatment of radiative transfer. The reason for this is very understandable: it is quite a numerical challenge to model multi-dimensional radiative transfer in protoplanetary disks that are very optically thick and actively accreting (and hence producing heat near the midplane). However, in recent years multi-dimensional radiative transfer tools, in particular those based on the Monte Carlo method, have improved dramatically in speed. Extreme optical depths are now not necessarily a problem anymore. We have developed a code, MCMMax, that can now handle extreme optical depths efficiently, even if many of the photon packages originate from the optically thickest regions of the disk (Min et al. 2009). In addition, we have adjusted the code to fully take into account the density dependend sublimation state of both the refractory as well as the icy material (Kama et al. 2009). It is therefore a right time to revisit the problem of the snow line with the full force of multi-dimensional radiative transfer and investigate (a) where is the ice line as a function of time and (b) what is the gas and solid surface density distribution in the disk as a function of time. This is the goal of this paper.

2. Model

Our model combines the radiative transfer equipment and basic disk model setup described in Min et al. (2009). We add here: a self-consistent and automatic computation of the radial structure, viscous heating of the disk by accretion and a self-consistent treatment of water ice and dust evaporation.

Right at this point we already note that we will replace ‘time’ \( t \) with ‘accretion rate’ \( \dot{M} \) and study only the inner 20 AU of the disk. The reason is that the long term evolution of protoplanetary disks depends very strongly on processes such as photoevaporation of the outer disk regions (see e.g. Gorti and Hollenbach 2009). The inner disk regions, say inward of 20 AU, are regulated entirely by the ‘feeding’ of matter by
the outer disk. How this feeding depends on time requires a detailed study of the complex processes happening at large radii. But given a recipe of feeding, i.e. an accretion rate as a function of time $\dot{M}(t)$, the inner disk can be assumed to be entirely determined. In particular if we avoid issues such as ionization degree and possible instabilities in the disk (see e.g. [Hartmann et al., 2006] and assume that no mass pile-up can happen anywhere in the inner disk (this assumption must be relaxed in future work), then the stationary disk structure inward of 20 AU is entirely determined by the stellar parameters and the accretion rate $\dot{M}$. For any given star our model thus becomes a 1-parameter model. So instead of having $t$ as our parameter, we will have $\dot{M}$ as our parameter. Generally, a high mass accretion rate corresponds to early times in the evolution of the solar nebula, while low mass accretion rates correspond to later times. Alternatively, in the episodic accretion scenario, low mass accretion rates represent the typical state, which is punctuated by relatively short episodes with high mass accretion rates.

The main improvement of our study over earlier work is that we treat the flow of radiation through the disk in a fully self-consistent manner, assuming that dust opacities are everywhere dominant over the gas opacities. Our model is axially symmetric, i.e. all variables depend on radial coordinate $r$ and vertical coordinate $z$, but do not depend on longitudinal coordinate $\phi$. The movement of photons is, however, fully 3-D, i.e. a photon can move also outside of the $r, z$-plane and has all the 3-D freedom of motion. In our Monte Carlo scheme (based on Bjorkman and Wood, 2001) we split the total input luminosity $L$ into $N$ 'photon packages', each carrying a luminosity $L_{\text{package}} = L/N$. The total luminosity includes the stellar luminosity $L_\star$ and the accretional heating luminosity $L_{\text{visc}}$, i.e. $L = L_\star + L_{\text{visc}}$. As these photon packages travel through the disk they update the local temperature of the dust according to the scheme by Bjorkman and Wood (2001). Since our model is axisymmetric, these updates have to only be done in 2-D, i.e. the temperature only depends on (and will only be updated in) $r, z$, no matter what is the longitudinal location $\phi$ of the photon package. So while the radiative transfer is 3-D, the resulting structure is 2-D.

The main difference between models doing radiative transfer in 1+1D and our model doing full 3D is in the radial energy diffusion. In the 1+1D model the energy from the star is intercepted by the disk and further diffusion is assumed to be vertical. This approximation is expected to be accurate for the regions of the disk close to the midplane where the very high densities do not allow much radial diffusion. However, radial energy diffusion is important in the surface layers and, more importantly, in the innermost regions of the disk, directly illuminated by the central star. Our model naturally takes this into account and allows us to see in which parts of the disk the vertical diffusion approximation is valid. Also, 3D radiative transfer is the only way to properly treat the complex shape of the dust condensation front close to the star (see also Kama et al., 2009).

For the central star in our model we take the Sun, i.e. the mass $M_\star = M_\odot$, the luminosity $L_\star = L_\odot$, and the stellar radius $R_\star = R_\odot$. The spectrum of the Sun is approximated by a blackbody of 5777 K. Although the parameters of the early Sun were likely different from this, we do not expect much difference on the results in this paper.
2.1. Treatment of viscous accretion

We assume that viscous heating is described by the so-called $\alpha$-disk model (Shakura and Sunyaev, 1973). In this model it is assumed that the energy produced locally by viscous heating is proportional to the local gas pressure through the proportionality constant $\alpha$. The total energy locally created per unit volume is then given by

$$\Gamma_{\text{viscous}} = \frac{9}{4} \alpha P(z) \Omega(R),$$

where $P(z)$ is the gas pressure, and $\Omega(R)$ is the Keplerian angular velocity at radius $R$ from the central star. If we define a uniform accretion rate $\dot{M}$ throughout the disk we have for the total energy released per unit surface area of the disk per unit time at a radius $R$ from the Sun is

$$F_{\text{viscous}} = \frac{3GM_* \dot{M}}{4\pi R^3} \left[1 - \left(\frac{R_*}{R}\right)^{1/2}\right],$$

where $G$ is the gravitational constant, and $M_*$ and $R_*$ are the mass and radius of the central star. Combining Eqs. 4 and 5 we can solve for the surface density of the disk for any given value of $\alpha$. In this paper we will consider two values of $\alpha$, 0.1 and 0.01. From Eq. 5 it is clear that the total amount of energy released is independent of the structure of the disk. This speeds up convergence since there is no feedback from the disk structure on the total energy that is released.

We basically treat the energy released by viscous stress in the same way as the energy released by the central star. Thus a fraction of the photon packages released in the Monte Carlo radiative transfer process is now released from the inner regions of the disk. Where previous studies that take into account radiation from viscous stress rely on assumptions on the radiative transfer in the optically thick inner regions, we treat the radiative transfer fully consistent in the framework of Monte Carlo radiative transfer as outlined by Min et al. (2009). The total accretional heating luminosity of the disk is given by the integral of Eq. 5 from the stellar surface, $R_*$, to the outer radius, $R_{\text{out}}$. Assuming $R_{\text{out}} \gg R_*$ this leads to

$$L_{\text{visc}} = \frac{1}{2} \frac{GM_* \dot{M}}{R_*}.$$

Note that a large fraction of this energy is most likely released inside the dust sublimation radius and thus can easily escape the disk. This energy does not influence the dust temperature and is thus not taken into account in our computation. In our method the energy produced through viscous heating is directly deposited from the gas onto the dust grains. This assumes a strong coupling between gas and dust, which is an appropriate assumption in the regions where the densities of both gas and dust are high. In the regions where a large fraction of the dust is evaporated due to high temperatures, we limit the amount of energy deposited on the grains by viscous heating by taking into account the number of collisions between the gas and the dust and assuming a gas molecule can at maximum transfer its total kinetic energy onto the grain in such a collision. This limits the total energy that can be transferred to the dust per unit time. We find that indeed in the optically thick regions this limit exceeds the available energy.
Table 1: Parameters of the three different dust models considered. The composition is given as mass fractions. The gas to solid ratio is the ratio between the mass in gas and the total available mass to form solids. $F_{\text{ice}}$ is the fractional increase in the mass available to form solids when water ice can condense. The value of the Rosseland mean opacity, $\kappa_R$, is taken at 160 K is computed considering all dust species including water ice.

<table>
<thead>
<tr>
<th>Model</th>
<th>Silicates</th>
<th>FeS</th>
<th>C-dust</th>
<th>Water ice</th>
<th>Gas to solid ratio</th>
<th>$F_{\text{ice}}$</th>
<th>$\kappa_R$ [cm$^2$/g]</th>
</tr>
</thead>
<tbody>
<tr>
<td>only CO</td>
<td>47%</td>
<td>14%</td>
<td>-</td>
<td>39%</td>
<td>97</td>
<td>1.63</td>
<td>572</td>
</tr>
<tr>
<td>($w = 0$)</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>only C-dust</td>
<td>24%</td>
<td>7%</td>
<td>20%</td>
<td>49%</td>
<td>49</td>
<td>1.97</td>
<td>544</td>
</tr>
<tr>
<td>($w = 1$)</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>mixture</td>
<td>32%</td>
<td>10%</td>
<td>13%</td>
<td>45%</td>
<td>66</td>
<td>1.84</td>
<td>508</td>
</tr>
<tr>
<td>($w = 0.5$)</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

by many orders of magnitude. However, in the inner regions, this limit prevents the small amount of dust grains available to be overloaded with accretional energy from the dense gas surrounding it. The dust grains are assumed to radiate the energy they absorbed with a wavelength distribution according to the local dust temperature after which the regular procedure of Monte Carlo radiative transfer is proceeded.

In the dusty regions of the disk the optical depth from the midplane, where most energy is released, to the outside is extremely large. Therefore, a photon package undergoes a large number of interaction steps before it leaves the system (on the order of several million interactions). If all these interaction steps would be computed separately, the computation time required would be on the order of a minute per photon package, putting strong constraints on the number of photon packages that can be used. Min et al. (2009) introduced an analytical solution for the random walk regime of the radiative transfer that can be applied in the optically thick regions and reduces computation times by orders of magnitude. This algorithm therefore allows the use of a large number of photon packages and thus to obtain proper statistics everywhere.

2.2. The dust and ice composition and opacities

In order to estimate the abundances of the various dust components we use here the Solar abundances of the elements as determined by Grevesse and Sauval (1998). The dust mixture is then constructed as follows:

- Since iron sulfide is the dominant form of (solid) sulfur in meteorites in the solar system it is first assumed that all the available sulfur is in the form of FeS. This takes approximately half the available number of Fe-atoms.

- Then we assume all Mg and Si go into silicates together with the remaining Fe-atoms. This creates silicates with an average composition of roughly MgFe$_{0.5}$SiO$_{3.5}$. We assume this to be in the form of amorphous silicates. This takes approximately 15% of the available O-atoms.
Next we put a fraction $w$ of the available C-atoms in the form of carbonaceous dust grains. The remaining C-atoms are assumed to form CO. This takes away up to $\sim 50\%$ of the available number of O-atoms depending on the value of $w$.

The remaining O-atoms are assumed to form water ice everywhere in the disk where the temperature allows it.

From the above procedure we also can compute the gas to dust ratio. We find that the resulting gas-to-dust ration can deviate up to a factor of 2 from the canonical value 100. The question remains what the most realistic value of the carbon abundance in dust, $w$, is. We take as a standard the value $w = 0.5$, based on in-situ measurements of the dust in comet Halley (Geiss, 1987). However, the carbon fraction in dust could have been higher in the early Solar nebula since the interstellar solid carbon abundance is much higher. Also, in the inner regions, at temperatures above $\sim 1000$ K, carbon grains will be destroyed by combustion so lower values of $w$ are to be expected. The dust/ice/gas mixture resulting from this procedure for different choices of $w$ are summarized in Table I.

In order to compute the opacities of the dust species as a function of wavelength we need to know the refractive indices of the grains. For the silicates we use the measurements by Dorschner et al. (1995) and Henning and Stognienko (1996), for the iron sulfide we take the laboratory data from Begemann et al. (1994), for the carboneous material we take the measurements from Preibisch et al. (1993), and for the water ice we take the data from Warren (1984). In order to convert the refractive index data to opacities we use the method by Min et al. (2005) to simulate irregularly shaped particles. The particle sizes are assumed to follow the so-called MRN distribution (Mathis et al., 1977). We thus assume that no significant grain growth has occurred with respect to the grains in the interstellar medium. We will come back to this assumption in the discussion of the results.

2.3. Treatment of evaporation and condensation

The temperature at which a species evaporates depends on the vapor pressure of the material. We treat the evaporation and condensation of the dust species following Kama et al. (2009). In order to converge the structure of the inner regions Kama et al. (2009) had to put constraints on the vertical and radial gradients of the fraction of the material in the gas phase. This is to avoid instabilities caused by the sudden jump in opacity at the inner edge of the disk. Since the ice is located in the inner regions of the disk, the ill conditioned convergence is no issue here. Therefore, for the water ice we do not put any constraints on the gradients of the condensed fraction. It turns out that convergence is reached relatively fast and no instabilities like those reported by Kama et al. (2009) are found for the ice sublimation zone. Although the density structure does not change significantly anymore after $\sim 30$ iterations, we ran all models up to 100 iterations to make sure also the slowly changing structures are converged.

To simplify the evaporation of the refractory dust species we take the evaporation parameters for amorphous olivine for the entire dust mixture of silicates, iron sulfide and carboneous material. For the iron sulfide it might be argued that this is not a correct assumption since the evaporation temperature of FeS is around 680 K. However,
we assume that above this temperature the iron condenses as metallic iron, contributing similarly to the opacity of the grains. The evaporation characteristics of both ice and olivine are based on sublimation data from Pollack et al. (1994).

2.4. Convection

In the implementation of radiative transfer and the computation of heat balance in the disk we employ, the effect of convection is not taken into account. However, we will show that convection is an important way of transporting energy generated in the midplane of the disk by accretion to the surface. Therefore, we have implemented a rough treatment of convection in the radiative transfer code to test the effects on our results. In the computations below, convection is not taken into account unless stated otherwise.

Convection becomes an important way of transporting heat when the radiative temperature gradient becomes too large. According to the Schwarzschild criterion for convection this happens when

$$\frac{d \ln T}{d \ln P} > \nabla_{\text{ad}},$$

where $T$ and $P$ are the gas density and pressure, and $\nabla_{\text{ad}}$ is the adiabatic limit which we take to be that of a diatomic gas: $\nabla_{\text{ad}} = 2/7$. Our rough implementation to correct the obtained temperature structure for convection is simply as follows. We start at the surface of the disk, where the temperature gradient is very small and work our way to the midplane. At each location the criterion according to Eq. 7 is determined. If the temperature gradient is too steep the temperature gradient is adjusted such that...
Figure 2: Same as Fig. 1 but for the case where all carbon is in the form of solid carbon.

Eq. (7) becomes an equality. In some cases this causes the temperature structure close to midplane to be adjusted such that the midplane temperature is reduced. This way we can model the effects of cooling of the midplane by convection although we should stay aware of the approximate nature of the implementation.

Our implementation of convection is very similar to the one by D’Alessio et al. (1998) with the difference that we do not consider the turbulent flux, and we take the convection efficiency, their parameter $\zeta$, to be unity. In the relatively less dense regions of the disk transfer of energy through radiation dominates and the criterion Eq. (7) is fulfilled. In this case we do not alter the temperature structure. However, in those regions of the disk where radiative transfer cannot fulfill Eq. (7) we do adjust the temperature structure to be adiabatic. This means that convection becomes important only in the more massive disks where radiative energy transfer is difficult.

3. Results

3.1. Location of the snow line as a function of $\dot{M}$

In Fig. 1 we display the locations in the disk where water ice can exist for various values of the mass accretion rate. It is clear that the location of the snow-line depends strongly on the value of $\dot{M}$. This is caused by two main effects. The first is that when $\dot{M}$ is increased, the energy released by viscous stress is increased. Secondly, a larger value of $\dot{M}$ causes the disk to be more massive (see also Fig. 4) and thus the greenhouse effect, keeping the energy at the midplane, is much stronger (as explained below, see also Eq. (8)) and the midplane temperature is increased further.

Going vertically up in the disk, the effect from viscous heating becomes less important and the snow line moves in significantly. This continues up to the point where radiation from the Sun becomes the dominant energy source and causes the snow line
to move out again. The fact that the location of the snow line in the upper, optically
thin region of the disk is not constant is caused by the pressure dependent evaporation
temperature of the ice in combination with a decreasing density when going higher up
in the disk. This shape of the ice region is consistent with the findings from [Davis
(2005a) but the snow line in our computations lies significantly further out. The rea-
son for this is most likely predominantly caused by the different opacity of the grains
which is assumed. Davis (2005a) use for the opacity the analytical expression from
Ruden and Lin (1986). As outlined below the most important parameter determining
the midplane temperature is the Rosseland mean opacity at 160 K divided by the gas to
solid ratio, since this determines how efficiently the energy created by viscous stress is
kept in the midplane at the location of the ice evaporation front. In the study by Davis
(2005a) this parameter is $\kappa_R/f = 2.7 \text{ cm}^2/\text{g}$, while in our study, taking a realistic dust
grain model, we find that $\kappa_R/f$ lies in the range 5.9 to 11.1 cm$^2$/g (see Table 1). Thus in
our model the energy generated by viscous heating is more efficiently captured in the
midplane regions.

In Figs. 2 and 3 we show the extreme cases of the range of dust parameters we
studied. Fig. 2 shows the case where all carbon atoms are locked into solid carbon, i.e.
the opacity is at its maximum. In this case the greenhouse effect, heating the midplane
efficiently is at its maximum. The other extreme, presented in Fig. 3, shows the case
where all carbon is in CO, i.e. the opacity is minimum, and the turbulence is very high,
$\alpha = 0.1$. This high value of $\alpha$ implies that one can get the same mass accretion rate with
a smaller surface density. Thus for a given value of $\dot{M}$ the surface density drops and
the greenhouse effect is smaller. In this case the energy can easily escape the midplane
and the midplane temperature is low leading to a snow line close to the star.
3.2. Analytical estimate for the location of the snow line in the midplane

We take the analytic estimate of the midplane temperature from [Hubeny, 1990]. This only includes viscous heating in the midplane and results in a midplane temperature, assuming high optical depths, of

\[ T^4 = \frac{9\Sigma_{\text{gas}}\kappa_R G M_* \dot{M}}{128\pi \alpha R^3 f}, \]  

(8)

where \( \kappa_R \) is the Rosseland mean opacity of the gas/ice mixture and \( f \) is the gas to solid ratio such that the total optical depth through the disk is equal to \( \tau_{\text{tot}} = \kappa_R \Sigma_{\text{gas}} / f = \kappa_R \Sigma_{\text{solid}} \). The values of \( \kappa_R \) and \( f \) for the models we computed are listed in Table 1. The surface density of the gas, \( \Sigma_{\text{gas}} \), can be expressed in terms of \( T \) and \( \alpha \) using Eqs. (4) and (5). The integral of Eq. (4) in vertical direction is equal to the total energy given in Eq. (5). Assuming \( R \gg R_* \) and taking the temperature to be the midplane temperature at each height in the disk we get (see also [Dullemond et al., 2007])

\[ \Sigma_{\text{gas}} = \frac{\mu m_p \dot{M}}{3\pi \alpha k_B T} \sqrt{\frac{G M_*}{R^3}}, \]  

(9)

where \( \mu \) is the average molecular mass and \( m_p \) is the proton mass. Substituting Eq. (9) into Eq. (8) we get

\[ T^5 = \frac{3\mu m_p M^2 \kappa_R}{128\pi^2 \alpha k_B \sigma f} \left[ \frac{G M_*}{R^3} \right]^{3/2}. \]  

(10)

Taking the above we find that the snow line is located at

\[ R_{\text{ice}} = \left[ \frac{C \cdot \dot{M}^2 \kappa_R}{f \alpha} \right]^{1/3}, \]  

(11)

where \( \kappa_R \) is the Rosseland mean opacity of the gas/ice mixture computed at 160 K (see Table 1), and \( C \) is given by

\[ C = \frac{3\mu m_p (G M_*)^{3/2}}{128\pi^2 \alpha k_B \sigma T_{\text{ice}}^5} = 1.7 \cdot 10^{22} \text{ cm}^{3/2} \text{ s}^{-2}/\text{g}. \]  

(12)

We find that this formula agrees very well with our findings for the range of parameters considered here (within a few percent). The formula becomes less accurate when the snow line is close to the star, and clearly the equation fails when the radiation of the Sun is the dominant energy source in the midplane. However, this is never the case in the models considered here. For easy computation, the value of \( C = 3.5 \cdot 10^{14} \) can be used in Eq. (11) when \( \dot{M} \) is given in \( M_\odot/\text{yr} \) and \( \kappa_R \) is given in \( \text{cm}^2/\text{g} \) to give the location of the snow line \( R_{\text{ice}} \) in AU.

Using Eq. (11) provides a way to scale our results to those obtained by [Davis (2005a)] taking into account the correct values of \( \kappa_R \), \( f \), and \( \dot{M} \). Doing this we find that the agreement is excellent, and thus the differences are indeed due to the different values of \( \kappa_R / f \). We also compared our results to those computed by [Garaud and Lin (2007)]. The values for the location of the snowline as computed using Eq. (11) are compared to
\begin{table}[h]
\centering
\begin{tabular}{|c|c|c|c|}
\hline
$\dot{M}$ [M$_{\odot}$/yr] & $\kappa_R/(f\alpha)$ [cm$^2$/g] & $R_{\text{ice}}$ [AU] (Eq. (11)) & $R_{\text{ice}}$ [AU] (numerical) \\
\hline
\textit{(Davis, 2005a)} & & & \\
8.06 $\cdot$ 10$^{-6}$ & $\sim$ 270 & 32.3 & $\sim$ 37 \\
1.75 $\cdot$ 10$^{-6}$ & $\sim$ 270 & 16.4 & $\sim$ 15.5 \\
1.13 $\cdot$ 10$^{-7}$ & $\sim$ 270 & 4.8 & $\sim$ 3.6 \\
2.53 $\cdot$ 10$^{-10}$ & $\sim$ 270 & 0.3 & $\sim$ 0.6 \\
\hline
\textit{(Garaud and Lin, 2007)} & & & \\
1 $\cdot$ 10$^{-7}$ & 27.7 & 2.8 & $\sim$ 3.4 \\
1 $\cdot$ 10$^{-8}$ & 27.7 & 1.0 & $\sim$ 1.4 \\
1 $\cdot$ 10$^{-9}$ & 27.7 & 0.4 & $\sim$ 0.6 \\
\hline
This study ($w = 0.5, \alpha = 0.01$) & & & \\
1 $\cdot$ 10$^{-6}$ & 770 & 16.1 & $\sim$ 16 \\
1 $\cdot$ 10$^{-7}$ & 770 & 5.8 & $\sim$ 6 \\
1 $\cdot$ 10$^{-8}$ & 770 & 2.1 & $\sim$ 2.2 \\
1 $\cdot$ 10$^{-9}$ & 770 & 0.7 & $\sim$ 0.8 \\
\hline
This study ($w = 0, \alpha = 0.1$) & & & \\
1 $\cdot$ 10$^{-6}$ & 59 & 9.1 & $\sim$ 9 \\
1 $\cdot$ 10$^{-7}$ & 59 & 3.3 & $\sim$ 3.2 \\
1 $\cdot$ 10$^{-8}$ & 59 & 1.2 & $\sim$ 1.2 \\
1 $\cdot$ 10$^{-9}$ & 59 & 0.4 & $\sim$ 0.5 \\
\hline
\end{tabular}
\caption{A comparison of the location of the snow line as obtained using Eq. (11) and using numerical computations. The values from the studies by Davis (2005a) and Garaud and Lin (2007) are read from the figures. The value of $\kappa_R/(f\alpha)$ for the study by Davis (2005a) is an estimate since they use a different description for the viscosity which corresponds approximately to ours with $\alpha = 0.01$. For the study by Garaud and Lin (2007) we used their Eq. (2) to determine the value of $\kappa_R$ at 160 K. The values of $R_{\text{ice}}$ for the numerical computations of this study correspond to the location where 50% of the ice is condensed.}
\end{table}

Table 2: A comparison of the location of the snow line as obtained using Eq. (11) and using numerical computations. The values from the studies by Davis (2005a) and Garaud and Lin (2007) are read from the figures. The value of $\kappa_R/(f\alpha)$ for the study by Davis (2005a) is an estimate since they use a different description for the viscosity which corresponds approximately to ours with $\alpha = 0.01$. For the study by Garaud and Lin (2007) we used their Eq. (2) to determine the value of $\kappa_R$ at 160 K. The values of $R_{\text{ice}}$ for the numerical computations of this study correspond to the location where 50% of the ice is condensed.

those obtained by Davis (2005a) and Garaud and Lin (2007) in Table 2. Note that Davis (2005a) adopted a different description for the viscosity of the disk, which explains partly the differences. The overall agreement with the approximated and numerically computed values implies that estimates of the location of the snow line using approximate methods are in general correct when appropriate opacity values are used.

3.3. Solid surface density distribution as a function of $\dot{M}$

In Fig. 4 we show the surface density as a function of radius for different values of the mass accretion rate. Both the total surface density (top panel) as well as the surface density in solids (lower panel) are shown. The sudden drop in the total surface density at the location where the solids start to become important is an effect of the sudden
increase in opacity, causing a steep temperature gradient which translates into a steep surface density gradient through Eqs. 4 and 5. The surface density in solids is adjusted iteratively until the temperature nowhere in the disk exceeds the sublimation temperature. This means in the inner regions solid matter is removed until the greenhouse effect, keeping the energy locked in the midplane regions of the disk, drops to a level where the temperature is low enough. For very high accretion rates, a lot of matter has to be removed to let the large amount of energy produced in the midplane escape. This balance between condensation and evaporation results in a thermostat keeping the midplane temperature at the silicate evaporation temperature over a significant part of the inner disk. When the accretion rate is reduced, more matter can condense. However, at the same time reducing the mass accretion rate reduces the total surface density and thus the amount of matter that can condense. Both these effects play a role in determining the surface density in the solid phase.

At the radius of water ice sublimation we have a gradual increase in the surface density in the solid phase. We find that the increase in the solid phase due to water ice is \( F_{\text{ice}} = 1.84 \) for the case where half of the carbon atoms is in the solid phase. This increase is quite small compared to the \( F_{\text{ice}} = 4.2 \) taken currently in the literature. We have to take into account here that we have only considered water ice condensation. If we take in addition CH\(_4\) and NH\(_3\) ice, which are the most important ice species after water ice, we gain about a factor 1.5 (see e.g. Dodson-Robinson et al., 2009) so we get \( F_{\text{ice}} \sim 2.8 \). Note however, that these ices condense at lower temperatures, so somewhat further out than the water-ice condensation radius we compute here.

The surface density is in principle a function of the mass accretion rate, the value of \( \alpha \) and the opacity of the grains. However, in the innermost region the dust evaporates to reduce the greenhouse effect such that the midplane temperature drops below the dust sublimation temperature. In this regime the solid surface density is adjusted to reach a given vertical optical depth, \( \kappa_R \Sigma_{\text{gas}} / f \), roughly following Eq. (8). The maximum vertical optical depth that can exist before the midplane gets too hot is thus only a function of the total energy released in the midplane, and thereby only of \( M \) (see Eq. 5). The amount of mass that can exist in this region is robustly determined through the mass accretion rate and the opacity of the grains and is independent of the gas density or the description of the accretion mechanism. Therefore, in this region we can easily scale the solid surface density as a function of \( M \) we find in this study to different values of the opacity of the grains when, for example, they grow to larger sizes.

### 3.4. Effects of convection

When studying the temperature structure in the disk in detail we find that often the temperature gradient is steeper than the adiabatic temperature gradient (see Eq.7). In these cases the midplane will cool by convection. If we correct for the convective cooling in the way discussed in section 2.4, the midplane temperature decreases in some cases significantly. We find that this only plays a role when the midplane temperature is dominated by accretion and in addition, is most important when the temperature is above \( \sim 500 \) K. Thus, the location of the snow line is only mildly affected. For the case of \( M = 10^{-6} \) \( M_\odot / \text{yr} \) we find that due to the decreased midplane temperature the fraction of condensed solids can increase locally by a factor of 5 with respect to the case where
convection is ignored. For low accretion rates convection becomes unimportant at almost all radii and the disk cools through radiation (see also D'Alessio et al., 1998). The reason for the dependence of the importance of convection on the surface density of the disk is simply due to the fact that less massive disks can easily cool through radiation, so convection is not needed to lose the energy from the midplane. In the more massive disks it is very difficult for the radiation to escape from the midplane and convective cooling is much more efficient. Our treatment of convection is somewhat simplified and more advanced methods, also including turbulent fluxes, should be implemented in the future to get a clear picture of the exact effects.

In Fig. 5 we plot the surface density in the solid phase for computations with and without convective cooling of the midplane. It is clear that for the high accretion rates convective cooling is important to gain extra solids in the region where the terrestrial planets are (i.e. in the inner few AU). When convective cooling is ignored we find that there is never enough mass in the solids around 1 AU to create the Earth. However, with convective cooling we find that there is much more mass. Note that the amount of mass in the solids can be increased even further for the high accretion phases by growing the grains to somewhat larger sizes, thereby decreasing the opacity and thus the midplane temperature. We conclude that convective cooling is an important mechanism to take into account when modeling the solid density distribution in the early Solar nebula.

![Figure 4: The surface density as a function of radius for different values of the mass accretion rate. The top panel shows the total surface density, while the lower panel shows the surface density in solids. Blue solid lines are for $\dot{M} = 10^{-6} M_\odot/yr$, red dotted lines for $\dot{M} = 10^{-7} M_\odot/yr$, green dashed lines are for $\dot{M} = 10^{-8} M_\odot/yr$, and purple dashed-dotted lines are for $\dot{M} = 10^{-9} M_\odot/yr$. In all cases half of the available carbon atoms are in the solid phase and $\alpha = 0.01$.](image)
Figure 5: The surface density as a function of radius for different values of the mass accretion rate. The solid lines are with convective cooling included, the dotted lines are when only radiative cooling is used (same as in Fig 4). In all cases half of the available carbon atoms can be in the solid phase and $\alpha = 0.01$. In all panels the dashed lines gives the MMSN according to Eq. 2 with and without ice.

4. Discussion

4.1. Implications for planet formation

In this section we compare our results with constraints from the minimum mass solar nebula (MMSN). We will use here the 'classical' MMSN as mentioned above. There is still quite some debate on what the minimum mass is needed to form our solar system (see e.g. Davis 2005b, Desch 2007, Crida 2009), and most of these models find the need for a more massive solar nebula than the classical one we use here. Therefore, our choice of the MMSN is conservative, and the computations could be repeated to find stronger constraints by using other models. In addition, planet
formation might be more dynamic, i.e. planets could accrete mass from different parts of the disk through dynamic interactions.

According to the MMSN (Eq. 2) in order to form the Earth one needs a solid surface density of 7.1 g/cm$^2$ at 1 AU. As can be seen directly from Fig. 5, this surface density in solids is never achieved in our standard model. This is because for the high accretion models the midplane gets too hot, so solids have to evaporate to let the energy escape, while for the low accretion models there simply is not enough mass at all. However, as discussed above, if we take into account convective cooling of the midplane the amount of condensed solids can increase by a factor of 5. This way the surface density at 1 AU is only slightly below the requirement from the MMSN. Inside 1 AU the surface density in solids is never even close to the MMSN, which might pose serious problems for the formation scenario of Venus.

Jupiter is generally assumed to form outside the snow line. Thus, when Jupiter formed at its current position, we have as an additional constraint that $R_{\text{ice}} < 5.2$ AU. For this the mass accretion rate needs to be smaller than $\dot{M} < 8 \cdot 10^{-8} M_\odot/\text{yr}$. Since the midplane temperature at the snow line should be smaller than 160 K, convective cooling is not important, so we can use our standard model. At this mass accretion rate the solid surface density is approximately $\Sigma_{\text{solid}} \approx 2$ g/cm$^2$, which is slightly lower than the mass needed according to Eq. 2. However, typical giant planet formation models require up to 3 to 4 times the value of the MMSN (see e.g. Pollack et al., 1996).

We thus find that our computations put constraints on the scenario of the formation of the Solar system in the context of the MMSN. However, as we already noted above, the MMSN might be an optimistic estimate, and in reality a much larger mass could be required. A possible solution to still meet the requirements of planet formation in that case would be to lower the opacity of the grains significantly by, for example, growing them to larger sizes.

4.2. Dust properties in the hot midplane regions

We find that at high mass accretion rates there is a significant region in the disk where the midplane temperature is above 1000 K. This is out to ~1.7 AU in the case where $\dot{M} = 10^{-6} M_\odot/\text{yr}$, and out to ~0.6 AU in the case of $\dot{M} = 10^{-7} M_\odot/\text{yr}$. This is including convective cooling. Thus we find that a considerable region in the early Solar system had conditions where thermal processing of the solid state material was very efficient. Also, at these temperatures grain coagulation is influenced by sintering (see e.g. Poppe, 2003). Sintering is the process where the bonds between grains are molten together by the impact of grain collisions or by a long exposure to high temperatures. The structures formed under these conditions can be extremely open aggregates when the strong bonds created do not allow restructuring or compaction of the aggregates. Poppe (2003) compute that for aggregates composed of silica spheres exposure to temperatures around ~ 1250 K is needed for efficient sintering. This is comparable to the crystallization temperature of silica found by Fabian et al. (2000). Since both sintering and crystallization are processes that are caused by the atoms in the lattice wanting to move to the minimum energy state, it is not unlikely that the temperatures of these processes are similar. In that case sintering of Fe/Mg silicates will happen at slightly lower temperatures, i.e. around ~ 1000 K. The collision speeds in the midplane regions of the disk are too low to cause collisional sintering.
As mentioned before, grain growth might help to increasing the solid surface density by reducing the opacity of the particles and in this way decreasing the temperature in the disk. A low opacity reduces the greenhouse effect and thus allows more grains to condense before the midplane temperature gets too high. In fact, when the Rosseland mean opacity computed at the dust evaporation temperature of 1500 K reduces by a given factor, the maximum allowed surface density increases by this factor. In order to test to what sizes we have to grow the grains in order to significantly increase the solid surface density we performed computations of the Rosseland mean opacity at 1500 K for larger grain sizes. We increased the entire size distribution by a given factor $\gamma$ which means that we do not only grow large particles, but efficiently remove the small particles as well. Thus, where the MRN distribution we used so far runs from 0.005 $\mu$m to 0.25 $\mu$m, the new size distribution contains grains with sizes from $(\gamma \times 0.005)\mu$m to $(\gamma \times 0.25)\mu$m. For these larger grain sizes the effects of anisotropic scattering become significant. The argumentation on the maximum total optical depth were made under the assumption of isotropic scattering. When anisotropic scattering is important, the same argumentation can be made when scaling the scattering cross section by a factor $(1 - g)$ (see Ishimaru 1978) where $0 < g < 1$ is the anisotropy parameter. For values of $g$ close to unity, most radiation is scattered in forward direction which for radiative transfer purposes is effectively a strong reduction of the effect of scattering. Values of $g$ close to zero are more equivalent to isotropic scattering.

The solid curve in Fig. 6 shows the increase in the maximum allowed solid surface density as a function of the scale factor $\gamma$. For example, we see that by increasing the grain size distribution by a factor 1000 the maximum allowed surface density increases by a factor $\sim$44. This is because at these grain sizes the Rosseland mean opacity (corrected for anisotropic scattering) is reduced by a factor of $\sim$44. The dotted curve in Fig. 6 shows the same but now for very fluffy aggregates. The opacities here are computed using the Aggregate Polarizability Mixing Rule (APMR; see Min et al. 2008). It is clear that compared to the grain growth computed using compact grains (the solid curve) the fluffy aggregates have to grow to much larger sizes in order to reduce the opacity sufficiently to allow more solid mass to condense. If indeed sintering causes rigid bonds, prohibiting compaction of the formed aggregates, such fluffy aggregates are not an unlikely outcome of the aggregation process.

An interesting feature for the compact grains (and to a lesser extend for the fluffy aggregates) in Fig. 6 is that the opacity of the grains first increases, which causes more mass to evaporate. Only when the particles are sufficiently large the opacity decreases and the solid surface density can increase. This is caused by the increasing scattering efficiency when the grains grow from interstellar grain sizes to several microns in size. It implies that in order to form large particles, a barrier has to be taken. The complex behavior of grain growth under these conditions asks for a more detailed study. For the fluffy aggregates this effect is much weaker. This is because for these aggregated structures the scattering is heavily forward peaked which makes the effect of scattering on the radiative transfer very small.

Another possibility to decrease the opacity of the grains in the hot regions is to take away the species that contribute most to the opacity. We took a mixture of silicates, iron sulfide and carbonaceous grains. It is well known that iron sulfide is only stable at temperatures below 680 K (Pollack et al. 1994). This means that the iron sulfide
opacity source is most likely not there, though it can be argued that metallic iron will take over. Removing iron sulfide from the mixture we decrease the Rosseland mean opacity at 1500 K by a factor ∼1.2, so this would allow only somewhat more mass to form. The dominant species of opacity in our computations is carbon. If we were to remove all the carboneceous species but leave the iron sulfide we reduce the Rosseland mean opacity at 1500 K by a factor ∼3. Indeed, we find that the surface density in the inner region for our model without carboneceous solids is a factor of 3 higher than our standard model. Removing both the carboneceous species and the iron sulfide reduces the Rosseland mean opacity at 1500 K by a factor ∼20. Though we did not perform the exact radiative transfer computations we can deduce that this would allow ∼ 20 times more mass in silicate grains to condense in the hot inner regions of the disk. This is assuming no other species adding opacity, like for instance metallic iron, form in these parts of the disk. Since metallic iron is a very strong opacity source (even slightly stronger than iron sulfide) even a small amount would already increase the opacity again to approximately the same level as with iron sulfide and carbon present.

In conclusion we find that the solid surface density in the region controlled by evaporation due to viscous heating is completely determined by the opacity of the grains and thus heavily depends on the local grain size and chemical composition.

4.3. Coagulation timescale in the thermostat region

Another important effect of the opacity-temperature control system is that the reduced number density of small dust grains will have a significant influence on the coagulation time scale in that region. The time scale at which coagulation can proceed depends strongly on the initial number density of grains at which the aggregation process has to start. Smaller particle densities increase the mean free path for particles and, consequently, the collision time for a dust particle to find another one to stick to and
grow. This effect is extended through the further growth phases. Dominik and Dullemond (2008) have shown that this effect extends the time scales for removing small grains from the disk by coagulation approximately inversely proportional to the initial number density. This means that that coagulation will be strongly slowed down in this region, and that the time scales over which opacity dominates the number of grains that can be present in the thermostat region might be determined by the mass accretion time scale of the disk rather than the coagulation time scale. We believe that this effect can have significant influence on coagulation and disk evolution models. Further studies on this issue are necessary and will be the subject of a follow-up paper.

5. Conclusions

We successfully computed the solid density structure and snow line in the early Solar nebula for various parameters of the mass accretion rate and grain composition. This was done using full 3-D radiative transfer taking into account both the energy released by viscous heating as well as radiation from the Sun. The density structure of both the gas and the solids are computed self consistently using an $\alpha$ description for the viscosity and detailed evaporation physics for the dust and ice. We find that:

- The location of the snow line is determined by the opacity of the grains and the mass accretion rate. We show that approximate radiative transfer methods to compute the location of the snow line are quite accurate.

- At high mass accretion rates the midplane temperature in the inner few AU can exceed the silicate evaporation temperature. This causes a balance between condensation and evaporation of matter acting like a thermostat keeping the midplane at the silicate evaporation temperature over a significant region. The solid mass in the inner regions is dramatically reduced by this effect and models computing the solid state surface density of the early Solar system should take this into account.

- For our standard set of parameters the surface density in the solid state in the region of the terrestrial planets is below the MMSN for all values of the mass accretion rate. This poses serious constraints on the scenarios of planet formation. It could point to a significantly lower opacity value at the time of planet formation, which can be achieved by increasing the average grain size or significantly change the chemical composition.

- At high accretion rates the increased midplane temperature of the disk due to viscous heating causes the midplane temperatures to be $> 1000\, \text{K}$ in a relatively large part of the disk. Under these conditions thermal processing of dust material as well as sintering of aggregates becomes important.

- We find that for the high mass accretion rates the effects of convective cooling of the midplane are important to take into account.
References


