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1 **Anaerobic oxidization of methane in a minerotrophic peatland: enrichment**  
2 **of nitrite-dependent methane-oxidizing bacteria**

3  
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16 **Abstract**

17 The importance of anaerobic oxidation of methane (AOM) as methane sink in freshwater  
18 systems is largely unexplored, particularly in peat ecosystems. Nitrite-dependent anaerobic  
19 methane oxidation (n-damo) was recently discovered and reported to be catalyzed by the  
20 bacterium '*Candidatus Methyloirabilis oxyfera*' that is affiliated with the NC10 phylum. So  
21 far, several *M. oxyfera* enrichment cultures have been obtained using a limited number of  
22 fresh water sediments or wastewater treatment sludge as inoculum. In this study, using stable  
23 isotope measurements and pore water profiles, we investigated the potential of n-damo in a  
24 minerotrophic peat land in the south of the Netherlands that is infiltrated by nitrate-rich  
25 ground water. Methane and nitrate profiles suggested that all methane produced was oxidized  
26 before reaching the oxic layer, and NC10 bacteria could be active in the transition zone where  
27 counter-gradients of methane and nitrate occur. Quantitative PCR showed high NC10  
28 bacterial cell numbers at this methane-nitrate transition zone. This soil section was used to  
29 enrich the prevalent NC10 bacteria in a continuous culture supplied with methane and nitrite  
30 at an *in situ* pH of 6.2. An enrichment of nitrite-reducing methanotrophic NC10 bacteria was  
31 successfully obtained. Phylogenetic analysis of retrieved 16S rRNA and *pmoA* genes showed  
32 that the enriched bacteria were very similar to the ones found *in situ*, and constituted a new  
33 branch of NC10 bacteria, with an identity percentage of less than 96% and 90% to the 16S  
34 rRNA and *pmoA* genes of *M. oxyfera*, respectively. The results of this study expand our  
35 knowledge of the diversity and distribution of NC10 bacteria in the environment, and  
36 highlight their potential contribution to nitrogen and methane cycles.

37

38

## 39 **Introduction**

40 Wetlands are the largest single source of methane with estimated emissions of 103 Tg per  
41 year, which account for about 20% to 40% of the global annual atmospheric methane flux (1,  
42 8, 19). It is estimated that about 50% of the methane produced in wetlands is consumed before  
43 it reaches the atmosphere; this significant microbial methane sink is usually considered to  
44 consist exclusively of aerobic methanotrophic bacteria, which degrade methane using oxygen  
45 as electron acceptor (2, 5, 19, 39). In ecosystems where oxygen is depleted but sufficient  
46 alternative electron acceptors, e.g. sulfate or nitrate are present, methane can also be  
47 converted anaerobically (25, 38). Anaerobic oxidation of methane (AOM) coupled to sulfate  
48 reduction is performed by a consortium of anaerobic methanotrophic archaea (ANME) and  
49 sulfate reducing bacteria (SRB) (25, 47). Its significance is well established for marine  
50 ecosystems, where it may consume more than 90% of the produced methane (39). In  
51 freshwater wetlands, and especially peatlands, electron acceptors are scarcer, with  
52 concentrations typically in the low  $\mu\text{M}$  range (37). Due to this reason, redox processes are  
53 mostly limited by electron acceptor supply, very dynamic and highly susceptible to alterations  
54 e.g. by influx of polluted groundwater and atmospheric deposition of nitrogen and sulfur  
55 species (18, 46). The influence of nitrogen pollution on methane oxidation is complex, and  
56 not all feedback loops are well understood (2, 3, 16, 29). In principal the role of the alternative  
57 electron acceptors nitrate and sulfate for diverting carbon fluxes away from methane  
58 production is better established, given that sulfate and nitrate reduction are  
59 thermodynamically more favorable than methanogenesis (17, 30, 31, 51). However, these  
60 alternative electron acceptors can in principle also enable methane oxidation (47, 54), but this  
61 topic has received only little attention with respect to methane cycling in peatlands (43).

62 In the meantime, for other freshwater ecosystems, more and more evidence about the  
63 occurrence of AOM coupled to sulfate (11, 40), iron(III) (42) and nitrate reduction (9, 38, 44,  
64 50) has become available. Whereas nothing is known about the microorganisms mediating  
65 ferric iron reduction with methane, for sulfate reduction a very similar consortium of  
66 methanotrophic Archaea and SRB as in marine ecosystems is hypothesized to be responsible  
67 (11, 40). Nitrate- or nitrite-dependent AOM (n-damo), when linked to organisms, was so far  
68 always found to be performed by one bacterial species affiliated to the NC10 phylum (9, 13).  
69 Genome sequencing, expression studies and physiological experiments indicated that this  
70 bacterium, then named *Candidatus Methyloirabilis oxyfera*, is an “intra-aerobic”  
71 methanotroph that produces its own oxygen from the dismutation of nitric oxide into  
72 dinitrogen gas and oxygen. The produced oxygen is then used for canonical aerobic methane  
73 oxidation starting with the methane monooxygenase enzyme complex (12). Although 16S  
74 rRNA sequences similar to *M. oxyfera*'s were found in various environments (14), so far n-  
75 damo enrichment cultures have only been obtained from two types of ecosystems: eutrophic  
76 freshwater sediments and wastewater treatment sludge. The dominant bacteria in all described  
77 cultures were closely related ( $\geq 97\%$  identity of the 16S rRNA gene sequence) to *M. oxyfera*  
78 (13, 14, 20, 33). Currently it is unclear, however, if *M. oxyfera*-related species are the only  
79 nitrite-dependent methane oxidizing bacteria; if methane oxidation is a general feature of  
80 NC10 phylum bacteria or limited to (close relatives of) *M. oxyfera*, and how important these  
81 bacteria are for methane cycling in various ecosystems.

82 In this paper, we studied a minerotrophic peatland infiltrated by nitrate-containing  
83 groundwater. At the sampling site, no methane emission was detectable. Porewater profiling  
84 revealed a nitrate-methane transition zone below the oxic layer that could provide an  
85 ecological niche for n-damo microorganisms. NC10 bacteria abundance in soil cores was  
86 analyzed using quantitative PCR, and the section with highest cell numbers of *M. oxyfera*,

87 coinciding with the methane-nitrate transition zone, was used as inoculum for the enrichment  
88 of n-damo bacteria. Mimicking field conditions as much as possible by using nitrite-amended  
89 peatland water in continuous cultivation, a new cluster of *M. oxyfera*-like bacteria was  
90 enriched.

91

## 92 **Materials and methods**

93 **Site description.** The Brunssummerheide peatland (50°55'39.63"N/5°59'50.73"E) is a small  
94 (15ha) spring fen located in an oligotrophic sandy valley fed by locally upwelling, weakly  
95 buffered nitrate-polluted groundwater. The peat layer is relatively thin (maximum 2.5 m) and  
96 vegetation is dominated by *Sphagnum spec.*, *Nartheceium ossifragum* and *Molinia caerulea*.  
97 At the sampling site, nitrate-enriched groundwater overflows the peatland surface and  
98 infiltrates into the peat layer.

99 **Porewater profile determination and soil sampling.** Nitrate and methane profiles were  
100 determined by measuring the concentrations in porewater samples collected using 5 cm  
101 ceramic cups (Eijkelkamp Agrisearch Equipment, Giesbeek, The Netherlands) connected to  
102 Teflon tubes. Porewater samples were obtained at least in duplicate from the depth of 20 cm  
103 to 220 cm at 5 or 10 cm intervals in December 2009 and June 2010. Porewater for methane  
104 analyses was collected in vacuumed anaerobic glass bottles (40 ml) prefilled with 5 g sodium  
105 chloride and sealed with butyl rubber stoppers. For chemical analyses, porewater was  
106 collected in 60 ml syringes. Samples were transported to the laboratory within two hours in a  
107 cooling box, and stored at 4 °C for maximum 14 days before analysis. Methane in the bottle  
108 headspace was measured after pressure equilibration with argon using gas chromatography as  
109 described previously (14). Nitrate was analyzed colorimetrically on a Traacs 800+ auto-  
110 analyzer as described previously (48). Redox potential measurements were performed by

111 gently pushing platinum electrodes into pre-drilled holes and allowing them to equilibrate.  
112 Stable readings were obtained after 30 min (15). Soil samples were obtained from 50 cm to  
113 130 cm depth with a Russian peat corer, sliced into 5-20 cm intervals in the field, immediately  
114 put into self-sealing plastic bags, and stored in air-tight bins with oxygen scavenger  
115 (Anaerogen, Oxoid, USA), then transported to the laboratory and stored anaerobically at 4 °C  
116 until further analysis.

117 **Incubation.** Initially, 200 ml soil slurry of the depth layers of 80-100 cm, 100-120 cm and  
118 120-135 cm (sampled in July 2009) were incubated in separate bottles (500 ml). Surface water  
119 from the peatland was collected and used for medium preparation after removal of particles  
120 by filtering through a hemo-filter (Hemoflow HF80S, Fresenius Medical Care, USA). The  
121 medium contained: 2 mM KHCO<sub>3</sub>, 0.2 mM Na<sup>15</sup>NO<sub>2</sub> (99.6% <sup>15</sup>N; Isotec, USA) and 0.5 mM  
122 NaNO<sub>3</sub>. The bottles were made anaerobic by 6 cycles of vacuuming and gassing with Ar/CO<sub>2</sub>  
123 (75:25), followed by 5 min of flushing with Ar/CO<sub>2</sub>. Then 10 ml <sup>13</sup>CH<sub>4</sub> was injected into the  
124 headspace (final concentration ca. 20%). The pH in the bottles was around 6.0 and the bottles  
125 were incubated at 25 °C, with gentle shaking at 100 rpm. The production of <sup>13</sup>CO<sub>2</sub> was  
126 measured by GC-MS in the headspace (see below).

127 After three months incubation, the bottle with strongest <sup>13</sup>CO<sub>2</sub> production was used as  
128 inoculum for continuous culturing in a 3 liter glass bioreactor (working volume 1.5 l;  
129 Applikon, Schiedam, The Netherlands) that was operated in sequencing batch mode to  
130 prevent biomass loss. One cycle constituted of 23 h of continuous supply of medium, 0.5 h  
131 settling, finally followed by 0.5 h discharging with a level-controlled pump. To keep the  
132 culture anaerobic, the reactor was continuously flushed with 20 ml min<sup>-1</sup> Ar/CO<sub>2</sub> (95:5) and 5  
133 ml min<sup>-1</sup> methane. The temperature was controlled at 25 °C and the pH at 6.0 to 6.2.  
134 Dissolved oxygen, temperature and pH in the reactor were monitored by respective

135 electrodes. Medium was prepared as described above, except using unlabeled nitrite. The  
136 nitrite concentration in the reactor was estimated daily with Merckoquant test strips (0-80 mg  
137  $\text{l}^{-1}$ ; Merck, Germany), and the concentration in the medium was slowly increased from 0.2  
138 mM to 2.5 mM dependent on the activity of the continuous culture. Nitrite concentrations in  
139 the reactor were kept below 20 mg  $\text{l}^{-1}$  (0.44 mM). The medium loading to the reactor was  
140 between 200 to 500 ml per day.

141 **Activity analysis.** Methane oxidizing activity in bottles was measured by determining the  
142 amount of  $^{13}\text{CO}_2$  produced from  $^{13}\text{CH}_4$  oxidation with GC-MS (Agilent 5975C inert MSD;  
143 Agilent, United States) as previously described (14). Activity in the reactor was tested in  
144 batch experiments with the whole culture. First medium supply was stopped and unlabeled  
145 nitrite was allowed to be depleted. The reactor was flushed with Ar- $\text{CO}_2$  (95:5) for 1 h while  
146 stirring, and checked for residual methane in the headspace. When undetectable, 0.2 mM  
147  $^{15}\text{NO}_2^-$  and 50ml  $^{13}\text{CH}_4$  were added. 20  $\mu\text{l}$  gas samples were taken every hour for  $^{13}\text{CO}_2$ ,  
148  $^{15,15}\text{N}_2$ ,  $^{15,14}\text{N}_2$  analysis. At the same time, 1 ml culture liquid was taken and centrifuged; the  
149 supernatant was kept at 4 °C for nitrite analysis. Nitrite concentrations were determined with  
150 colorimetric methods as described elsewhere (23). The influence of pH on activity was  
151 determined in batch incubations of 10 ml biomass in 40 ml serum bottles, buffered with MES  
152 (2-(N-morpholino) ethanesulfonate, 20 mM) to pH values between 5.9 and 6.7, and with  
153 MOPS (3-(N-morpholino) propanesulfonate, 20 mM) to pH values between 6.75 and 7.4  
154 (measured at the end of incubation).

155 **DNA isolation.** Total DNA from soil samples was isolated with the PowerSoil<sup>o,r</sup> DNA  
156 isolation kit (MO BIO Laboratories Inc., USA) according to the manufacturer's manual.  
157 Approximately 0.3 g homogenized soil was used for DNA isolation, and two independent  
158 isolations were carried out for each depth interval. DNA was eluted three times with pre-



159 warmed Milli-Q water from the column to ensure that the entire DNA had been collected.  
160 DNA in the third elution was undetectable by agarose gel electrophoresis ( $<0.2 \text{ ng } \mu\text{l}^{-1}$ ). DNA  
161 obtained from the same depth interval was pooled for qPCR analysis to minimize the  
162 influences from soil inhomogeneities. DNA from enrichment cultures was isolated with a  
163 method based on bead-beating and SDS lysis, as described previously (14). DNA quality was  
164 checked on agarose gel, and concentrations were measured in triplicate with NanoDrop (ND-  
165 1000, ISOGEN Life Science, The Netherlands).

166 **Quantitative PCR.** In order to quantify n-damo bacteria and all bacteria in different depths of  
167 the soil cores, quantitative PCR (qPCR) targeting the 16S rRNA gene was performed. To  
168 account for imperfect primer matching and known variability of results (14), two different  
169 primer pairs were used for each group. For NC10 phylum bacteria, primer pairs p1F & p1R  
170 and p2F & p2R, and for all bacteria, primer pairs 1100F & 1492R and 533F & 805R (Table 1)  
171 were applied. All q-PCR assays were performed according to the MIQE guidelines (Minimum  
172 Information for Publication of Quantitative Real-Time PCR Experiments) (4). qPCR  
173 experiments were carried out with the Bio-Rad IQTM 5 cycler real-time detection system  
174 using IQTM SYBR green Supermix (Bio-Rad, United States) in 25 $\mu\text{l}$  reaction volume as  
175 previously described (14), except using 65 °C for n-damo specific primer pairs and 58 °C for  
176 universal primer pairs as annealing temperature, which had been determined as most suitable  
177 for the present samples by gradient PCR. The qPCR reactions were carried out in 96 well  
178 plastic plates (Bio-Rad, United States) sealed with Opti-Seal Optical disposable adhesive  
179 (BIOplastics, The Netherlands). Fluorescence signals were obtained at 72 °C at the end of the  
180 elongation step of each cycle. PCR products obtained with n-damo specific and universal  
181 bacteria primer pairs were cloned and sequenced using the vector pGEM-T Easy (Promega,  
182 United states). The sequences retrieved were of the correct length (201 bp for p1F & p1R; 292  
183 bp for p2F & p2R; 291 bp for 515F & 805R and 410 bp for 1100F & 1492R), and the

184 obtained n-damo sequences were similar (>97.2% identity) to the sequence of *M. oxyfera*  
185 (accession no. FP565575). Standard curves for n-damo bacteria and general bacteria were  
186 constructed with plasmids containing corresponding inserts, taking into account the molecular  
187 mass of the plasmid including the insert, and the plasmid concentration. Plasmids copy  
188 numbers used as standard were between 30.7 to  $3.07 \times 10^8 \mu\text{l}^{-1}$  for NC10 bacteria, 86.9 to  
189  $8.69 \times 10^8 \mu\text{l}^{-1}$  for all bacteria. Two soil cores with partial overlap were analyzed. Both cores  
190 were sliced in sections between 5 and 10 cm in the field (see soil sampling and DNA  
191 isolation). In Fig. 1, each depth interval is represented by its average depth. DNA isolated  
192 from soil of 85-90 cm depth was used to test dilution effect; 10 times and 100 times had a  
193 maximum difference of 8.7%, compared with non-diluted ones. For NC10 bacteria non-  
194 diluted DNA was used as templates; but for primers targeting all bacteria, 100 times diluted  
195 DNA was used. PCR efficiencies calculated based on standards were between 90.6% and  
196 99.2%. Both standards and samples were run in triplicates. The copy numbers in samples  
197 were calculated based on comparison with the threshold cycle values of the standard curve,  
198 taking into account the dilution and the amount of total DNA obtained per gram soil.

199 **Phylogenetic analysis.** PCR was performed with DNA isolated from the soil layer used as an  
200 inoculum (80-100 cm depth), the enrichment culture after 3 months of incubation in bottles,  
201 and the continuous culture after 1 and 17 months of enrichment in the reactor. 16S rRNA  
202 sequences of n-damo bacteria were obtained with universal bacteria primer 8F or n-damo  
203 specific primer 193F in combination with n-damo specific primer 1043R (Table 1). PCR  
204 products of the correct size were ligated into the pGEM-T Easy cloning vector (Promega,  
205 United States) and amplified in *Escherichia coli* DH5 $\alpha$ . Plasmids were isolated from 10 to 15  
206 randomly selected white colonies per library using the GeneJet miniprep kit (Fermentas,  
207 Lithuania), and were sequenced at the DNA Diagnostics Center of Nijmegen University  
208 Medical Center. The sequences were aligned to reference sequences with the MUSCLE

209 algorithm. Phylogenetic trees were constructed with MEGA5 using the neighbor-joining  
210 method and the robustness of tree topology was tested by bootstrap analysis (1000 replicates).

211 With the same DNA samples also functional gene (particulate methane monooxygenase  
212 subunit A, *pmoA*) clone libraries were constructed. The particulate methane monooxygenase  
213 catalyzes the first step of methane oxidation and is well conserved in methane oxidizing  
214 bacteria, therefore *pmoA* is widely accepted as a marker gene for assessing diversity of  
215 aerobic and *M. oxyfera*-like anaerobic methanotrophs in the environment (34, 36). Two  
216 different forward primers targeting either most methanotrophs (A189b) or only close relatives  
217 of *M. oxyfera* (cmo182) were combined with a specific reverse primer (cmo682) (Table 1). A  
218 *pmoA* phylogenetic tree based on nucleotide sequences was constructed as described above.

219 **Fluorescence in situ hybridization.** On a monthly basis, 1.5 ml biomass was harvested from  
220 the reactor and forced through a 0.5 mm needle to break big cell aggregates. Then the sample  
221 was centrifuged and the pellet was washed twice with 1 ml 1×PBS, and fixed with  
222 paraformaldehyde on ice for three hours. Fluorescence in situ hybridization (FISH) was  
223 performed as previously described (13), using 40% formamide concentration. The following  
224 oligonucleotide probes were used: S-\*-DBACT-0193-a-A-18 and S-\*-DBACT-1027-a-A-18  
225 specific for n-damo bacteria (38) and a mixture of EUB I-III and V for most Bacteria (7).  
226 Images were acquired with a Zeiss Axioplan 2 epifluorescence microscope equipped with a  
227 CCD camera, together with the Axiovision software package (Zeiss, Germany).

228 **Nucleotide sequences accession numbers.** Representative 16S rRNA and *pmoA* gene  
229 sequences were deposited at the National Center for Biotechnology Information  
230 (<http://www.ncbi.nlm.nih.gov/>) with the accession numbers JX262153- JX262155 (*pmoA*)  
231 and JX262156-JX262161 (16S rRNA).

232

## 233 **Results**

234 **Porewater profiles.** Porewater depth profiles of the Brunssummerheide sampling location  
235 were determined on five occasions between July 2009 and May 2011, with an overall very  
236 similar pattern. Representative winter (December 2009) and summer (June 2010) profiles are  
237 shown in Fig. 1. Nitrate concentration decreased with depth, and became undetectable below  
238 100 cm. No methane was detected in the upper 80 cm, but methane gradually increased below  
239 the depth of 80 cm and reached the maximum concentration at around 120 centimeters (Fig.1  
240 A, B). Redox data indicated that the soil was completely anoxic below 50 cm depth, and  
241 living roots of vascular plants were not found below 60 cm depth. The maximum  
242 concentration of nitrate (0.6 mM) in June 2010 (Fig.1 B) was about 0.2 mM higher than that  
243 in December 2009 (Fig. 1. A), possibly due to relatively stronger evaporation of surface water  
244 and higher groundwater influx in summer. The maximum concentration of methane remained  
245 similar in both seasons, as well as the overall pattern: an opposing gradient at around 80-100  
246 cm depth.

247

248 **Quantifying abundance of NC10 bacteria in different soil depths.** Total bacterial and  
249 NC10 phylum abundance in different soil depths was determined in two overlapping cores by  
250 qPCR using primers targeting the 16S rRNA genes. The highest cell numbers ( $1.3\text{-}3.2 \times 10^7 \text{ g}^{-1}$   
251 wet soil) of NC10 bacteria were found at 80-85 cm depth (Fig. 1 C, D), coinciding with the  
252 concomitant decrease of methane and nitrate (Fig. 1 A, B), and a peak in abundance of NC10  
253 phylum-characteristic phospho-lipid fatty acids (Fig. 1 G; Kool et al, submitted). In contrast,  
254 total bacteria cell numbers, ranging from  $0.9$  to  $11.8 \times 10^8 \text{ cells g}^{-1}$  wet soil, did not show a  
255 depth-related pattern (Fig. 1 E, F).

256

257 **Enrichment and activity.** Nitrite dependent methane oxidizing activity was initially  
258 determined by measuring the fraction of  $^{13}\text{CO}_2$  in total  $\text{CO}_2$  after supply of  $^{13}\text{CH}_4$  and nitrite  
259 to three soil sections (80-100; 100-120; 120-135 cm). Despite the addition (and permanent  
260 presence) of nitrate (0.5 mM), all soil cores produced some methane in the first two weeks of  
261 incubation, but no methane oxidation could be detected (detection limit approximately 0.5  
262  $\text{nmol d}^{-1} \text{g}^{-1}$  soil). After about 3 month's incubation, the 80 - 100 cm section showed methane  
263 oxidation activity ( $9.0 \text{ nmol d}^{-1} \text{g}^{-1}$  soil, assessed as  $\text{CO}_2$  production), and an increase in this  
264 rate indicated microbial growth. This incubation (80-100 cm) was used as inoculum to start a  
265 sequencing batch reactor for the enrichment of the responsible microorganism. Over the first  
266 9 months of enrichment activity remained low with a nitrite reduction rate of about  $50 \mu\text{mol d}^{-1}$   
267  $\text{L}^{-1}$ , and then started to increase to about  $1.0 \text{ mmol (NO}_2^-) \text{ d}^{-1} \text{L}^{-1}$  in month 15. Batch tests  
268 and experience with previous NC10 bacteria enrichment cultures had indicated that nitrite was  
269 preferred over nitrate; consequentially the medium, prepared with *in situ* water was not only  
270 amended with nitrate, but also nitrite. To test the coupling of nitrite reduction to methane  
271 oxidations, both activities were tested in batch experiments after 10 months with  $^{15}\text{N}$ - and  $^{13}\text{C}$ -  
272 labelled substrates during the enrichment period (Fig. 2). Nitrite-N was completely recovered  
273 as nitrogen gas, concomitantly methane was fully oxidized to  $\text{CO}_2$ . The ratio of  $^{13}\text{CO}_2$  and  
274  $^{15,15}\text{N}_2$  production was 3:4.3, similar to the theoretical stoichiometry of 3:4 (38). An activity  
275 test at different pH values demonstrated that the culture preferred circumneutral conditions,  
276 but was active down to the lowest tested value of 5.9 (Fig. 3).

277

278 **FISH analysis of the enriched bacteria.** FISH was performed on biomass of the enrichment  
279 culture fixed every month, but no clear hybridization with NC10 specific probes was observed  
280 until after 8 months of medium supply. Even though small numbers of NC10 bacteria must

281 have been present, they remained undetectable at first due to strong autofluorescent background  
282 and hybridization inhibition, presumably caused by peat material. Starting at month 9, NC10  
283 cells could be detected (Fig. 4A). With the progression of incubation, both total cell numbers  
284 visualized by DAPI stain, and the percentage belonging to the NC10 phylum gradually  
285 increased (Fig. 4 B, C) and coincided with an increased activity of the culture. At month 14  
286 about 50%, and at month 19 more than 80% of the population did hybridize with the NC10  
287 specific probes (Fig. 4).

288 **16S rRNA and *pmoA* gene phylogenetic analysis.** *M. oxyfera*-related 16S rRNA and *pmoA*  
289 genes were successfully obtained from both inoculum soil and the enrichment after 1 or 17  
290 months of incubation. Long (>1000 bp) 16S rRNA sequences obtained with primer 8F  
291 (universal) and 1043R (NC10 specific) were used for phylogenetic analysis. Results showed  
292 that the 16S rRNA sequences belong to the group A of NC10 bacteria (14), forming a cluster  
293 (differences between 0.1% to 2.7%) with sequences retrieved from coal-tar contaminated  
294 aquifer (AF351214, AF351217, FJ810544) and lake Constance sediment (HQ906524,  
295 HQ906538) (9). These sequences share only 94.9% to 95.5% identity with *M. oxyfera* (Fig. 5  
296 A).

297 The phylogenetic analysis of the *pmoA* gene showed similar results. *pmoA* sequences from  
298 both soil and enrichment culture again cluster together with *pmoA* sequences retrieved from  
299 Lake Constance sediment (HQ906571, HQ906568, HQ906566) (9). These *pmoA* sequences  
300 had an identity with those of *M. oxyfera* of 86.2-90.9% on nucleotide level, but the 95.8-  
301 97.9% on amino acid level indicated functional conservation (Fig. 5 B). No significant  
302 difference could be observed between the inoculum and the 17 months old enrichment  
303 culture, indicating that no population shift within the NC10 phylum had occurred. Both the

304 16S rRNA and *pmoA* genes phylogenetic results suggested that a new cluster of NC bacteria  
305 had been enriched.

306

## 307 **Discussion**

308 The Brunssummerheide peatland is a spring fen in an oligotrophic sand valley fed by nitrate-  
309 polluted groundwater, and therefore contains nitrate concentrations in the upper peat layer  
310 which are uncommonly high for pristine peatlands (52). Also in contrast to many other  
311 peatlands (6, 24, 26, 27), methane was not detected in the upper 70-80 cm of the depth profile  
312 at 5 sampling occasions in different seasons from 2009 to 2011, even though methane was  
313 produced in the deep anoxic zone (below 100 cm, Fig. 1A, B). As roots of vascular plants do  
314 not reach that deep in the Brunssummerheide (maximum 60 cm), this suggested the existence  
315 of an anoxic methane sink in the peat, independent from oxygen and aerenchymal transport by  
316 roots, for which oxidized nitrogen species could serve as electron acceptor. The counter  
317 gradient of methane and nitrate at the depth of 80 cm may provide an ideal niche for, and may  
318 be at least partly created by the recently characterized n-damo bacteria. Targeting their 16S  
319 rRNA gene in DNA extracts from different depths confirmed this: Highest n-damo cell  
320 numbers (up to  $3.2 \times 10^7$  cells  $g^{-1}$  soil) and ratios (3 to 8% of total bacterial community) were  
321 observed at the depth of 80-90 cm (Fig. 1 C, D), coinciding with the methane-nitrate  
322 transition zone (Fig. 1 A, B). At this depth, also a peak in abundance of phospho-lipid fatty  
323 acids diagnostic for NC10 phylum bacteria was detected (Fig. 1 G; Kool et al., submitted).  
324 The n-damo cell number and lipid profiles also agreed with the finding that among soil  
325 samples from 80-100, 100-120 and 120-135 cm depth only the 80-100 cm sample showed  
326 anaerobic methane-oxidizing activity upon incubation. Despite the relatively high numbers of  
327 n-damo bacteria detected at a depth 80-90 cm, it took several months to obtain an enrichment

328 culture with measurable activity. Also detection by fluorescence *in situ* hybridization using  
329 NC10 phylum-specific probes, hampered by a strongly auto-fluorescent background from the  
330 organic-rich inoculum, was only possible after 9 months of continuous cultivation with  
331 constant supply of nitrite and methane. This may be due to the “dilution” of the naturally  
332 NC10 phylum-enriched soil layer with less active deeper layers (90-100 cm) in the inoculum,  
333 and a very low growth rate at the prevailing conditions, especially the pH (6.0 – 6.2). The pH  
334 optimum test showed that the NC10 phylum bacteria enriched in the continuous culture were  
335 only acidotolerant to a certain extent, not acidiphilic. They were active down to a pH below 6,  
336 but their physiological optimum was clearly higher, above 7 (Fig. 3). This is a prime example  
337 for the discrepancy between physiological and ecological optimum. In contrast to previous *M.*  
338 *oxyfera* enrichment cultures from neutral, eutrophic sediments (14), which had a similar  
339 optimum (around 7.5), but were not active at a pH below 7 (assessed under similar conditions,  
340 O. Rasigraf, MSc thesis, 2011, unpublished), a different ecotype was dominant in the more  
341 acidic and low nutrient environment. According to the species delineation of 97% identity of  
342 the 16S rRNA gene for bacteria in general and 93% of the *pmoA* gene diagnostic for  
343 methanotrophic bacteria (35), the NC10 phylum bacterium dominating the  
344 Brunssummerheide enrichment culture even constitutes a new species within the genus  
345 *Methylomirabilis*.

346 Like other NC10 enrichment cultures (14, 20, 33), the enrichment period was characterized by  
347 a long phase without measureable activity, followed by a period of slow, but exponential  
348 increase in nitrite consumption rate. In the present case, nitrite-reducing activity remained low  
349 for the first 9 months, and then started to increase to about 1.0 mmol (NO<sub>2</sub><sup>-</sup>) d<sup>-1</sup> L<sup>-1</sup> in month  
350 15. After this increase it was not possible to stimulate the growth of the culture further and a  
351 sort of stationary phase was reached similar to other enrichments of NC10 bacteria (14, 20,  
352 22). The doubling time of the Brunssummerheide *Methylomirabilis* strain was estimated to be



353 about two months, which is 4 to 8 fold lower than the values reported before (14). It is  
354 difficult to predict whether this reflects the growth rate under field conditions. On one hand  
355 some factors like a higher temperature (25 °C), the optimum temperature of methanotrophs in  
356 most peat soils (19) in contrast to 10-15 °C *in situ* and constant substrate supply may be  
357 beneficial, but other factors like stirring, use of surface- instead of porewater or a decrease in  
358 microbial partner communities may also be disadvantageous for growth in the laboratory.

359 However, once established, the methanotrophic community does not need to grow fast to  
360 constitute a relevant methane sink in the environment. According to previous estimations,  
361 *Methylomirabilis* cells in an enrichment culture have an activity of 0.1 to 0.4 fmol CH<sub>4</sub> cell<sup>-1</sup>  
362 d<sup>-1</sup> (14), indicating that the Brunssummerheide soil of 80 – 85 cm depth with about 1.3 to  
363 3.2×10<sup>7</sup> cells g<sup>-1</sup> soil may convert between 1.3 and 12.8 nmol CH<sub>4</sub> d<sup>-1</sup> g<sup>-1</sup> soil. This range is at  
364 the lower end of methane oxidation rates reported for aerobic methanotrophs (41) in wetlands,  
365 but apparently high enough to balance the methane diffusing upwards from deeper,  
366 methanogenic soil layers.

367 Nitrite is clearly the preferred electron acceptor of previously reported *M. oxyfera*  
368 enrichments (13, 14, 20, 38). When nitrite was depleted in the present *Methylomirabilis*  
369 enrichment culture, methane oxidizing activity in the presence of nitrate (1 mM) ceased; upon  
370 addition of fresh nitrite, methane consumption started again (data not shown), demonstrating  
371 that the methane-oxidizing activity of Brunssummerheide enrichment is also nitrite  
372 dependent. Although nitrite was also detected in the depth profile, its concentrations were  
373 much lower (max. 4.2 μM, mostly around the detection limit of the colorimetric method) than  
374 those of nitrate. There was no depth-related pattern, and values were not constant over time.  
375 The nitrite needed by n-damo bacteria active in the soil might be supplied by other  
376 microorganisms (e.g denitrifying bacteria) or *Methylomirabilis* itself converting nitrate to

377 nitrite using organic carbon compounds other than methane. This would explain why nitrate is  
378 sufficient as an electron acceptor for methane oxidation *in situ* and in the initial batch  
379 incubations, whereas after enrichment, concomitant with a relative loss of other bacteria and a  
380 degradation of labile organic carbon, this supply path is insufficient and nitrite addition  
381 becomes mandatory for methane oxidation.

382 The present study shows an additional, so far hardly investigated pathway linking the  
383 biogeochemical cycling of nitrogen and methane in peatlands. Given the world-wide  
384 increasing groundwater nitrate and atmospheric nitrogen loads (32, 45), this methane sink  
385 may become more relevant for mitigating the mobilization of carbon in the form of methane  
386 from wetlands in the future.

387

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Table 1 Primer pairs used for qPCR analysis and clone libraries construction in this study.

	Forward primer			Reverse primer			Annealing temp. (°C)	Target
	Designation	Sequence (5' - 3')	Reference	Designation	Sequence (5' - 3')	Reference		
qPCR	p1F	GGGCTTGACATCCCACGAACCTG	(14)	p1R	CGCCTTCCTCCAGCTTGACGC	(14)	65	NC10 bact. 16S rRNA
	p2F	GGGGAAGTCCAGCGTCAAG	(14)	p2R	CTCAGCGACTTCGAGTACAG	(14)	65	NC10 bact. 16S rRNA
	533F	GTGCCAGCMGCCGCGGTAA	(49)	805R	GACTACCAGGGTATCTAATC	(28)	58	All bact. 16S rRNA
	1100F	YAACGAGCGCAACCC	(10)	1492R	GGTTACCTTGTTACGACTT	(53)	58	All bact. 16S rRNA
Clone library	8F	AGAGTTTGATYMTGGCTCAG	(21)					
	193F	GACCAAAGGGGGCGAGCG	(14)	1043R	TCT CCA CGC TCC CTT GCG	(14)	55-65	NC10 bact. 16S rRNA
	A189b	GGNGACTGGGACTTYTGG	(34)					
	cmo182	TCACGTTGACGCCGATCC	(34)	cmo682	AAAYCCGGCRAAGAACGA	(34)	55-65	NC10 bact. <i>pmoA</i>

396 **Figure captions**

397 **Fig. 1** Depth profiles of the Brunssummerheide peatland. (A, B), Nitrate (filled square) and  
398 methane (open circle) concentrations in porewater sampled in December 2009 (A) and in June  
399 2010 (B). (C-F) Bacterial cell numbers (cells g<sup>-1</sup> wet soil) as assessed by qPCR on DNA  
400 extracted from two overlapping soil cores from 51 to 102 cm (open circles), and from 77 to  
401 127 cm (open triangles). NC10 bacteria abundance was determined with primer pairs p1F &  
402 p1R (C), and p2F & p2R (D). Total bacterial abundance was determined with primer pair 535F  
403 & 805R (E), and 1100F & 1492R (F). (G) Relative abundance of the phospho-lipid fatty  
404 acids 10-methyl-hexadecanoic acid (10MeC16:0, open diamonds) and 10-methyl-  
405 hexadecanoic acid (10MeC16:1Δ7, multiplied by 3, closed squares) diagnostic of NC10  
406 bacteria (data from Kool et al., [24A]).

407

408 **Fig. 2** Activity test of the enrichment culture at month 10 with <sup>15</sup>NO<sub>2</sub><sup>-</sup> and <sup>13</sup>CH<sub>4</sub>. Nitrite  
409 (filled circle) was consumed, <sup>15,15</sup>N<sub>2</sub> (filled square), <sup>14,15</sup>N<sub>2</sub> and <sup>13</sup>CO<sub>2</sub> (filled triangle) were  
410 produced. The <sup>13</sup>CO<sub>2</sub> production rate was 20.2 μmol d<sup>-1</sup>, and the rate of <sup>15,15</sup>N<sub>2</sub> production  
411 was 29.0 μmol d<sup>-1</sup>.

412

413 **Fig. 3** Methane-oxidizing activity of the n-damo enrichment culture incubated at different pH  
414 values.

415

416 **Fig. 4** Fluorescence in situ hybridization of the enrichment culture at different times of  
417 incubation. A: month 9; B: month 14; C: month 19. NC10 bacteria appear in pink, due to co-  
418 hybridization of NC10 bacteria specific probes 193-Cy3 and 1027-Cy3 (red) and a mixture of  
419 probes EUBI-III, IV-Cy5 (light blue) for most eubacteria and DAPI (dark blue). (Scale bars:  
420 5µm).

421

422 **Fig. 5** Phylogenetic trees of the 16S rRNA (A) and the *pmoA* genes (B, including *amoA* and  
423 *pxmA* sequences) of the enrichment culture. The trees were calculated in Mega5 using the  
424 neighbor-joining method. Bootstrap support values (1000 replicates) greater than 50% are  
425 indicated at the nodes. The sequences obtained in this study from inoculum soil and  
426 enrichment after 1 or 17 months of incubation are shown in bold. **References**

- 427 1. Bastviken, D., L. J. Tranvik, J. A. Downing, P. M. Crill, and A. Enrich-Prast. 2011.  
428 Freshwater Methane Emissions Offset the Continental Carbon Sink. *Science* 331:50.
- 429 2. Bodelier, P. L. E. 2011. Interactions between nitrogenous fertilizers and methane cycling in  
430 wetland and upland soils. *Current Opinion in Environmental Sustainability* 3:379-388.
- 431 3. Bodelier, P. L. E., and H. J. Laanbroek. 2004. Nitrogen as a regulatory factor of methane  
432 oxidation in soils and sediments. *Fems Microbiology Ecology* 47:265-277.
- 433 4. Bustin, S. A., V. Benes, J. A. Garson, J. Hellemans, J. Huggett, M. Kubista, R. Mueller, T.  
434 Nolan, M. W. Pfaffl, G. L. Shipley, J. Vandesompele, and C. T. Wittwer. 2009. The MIQE

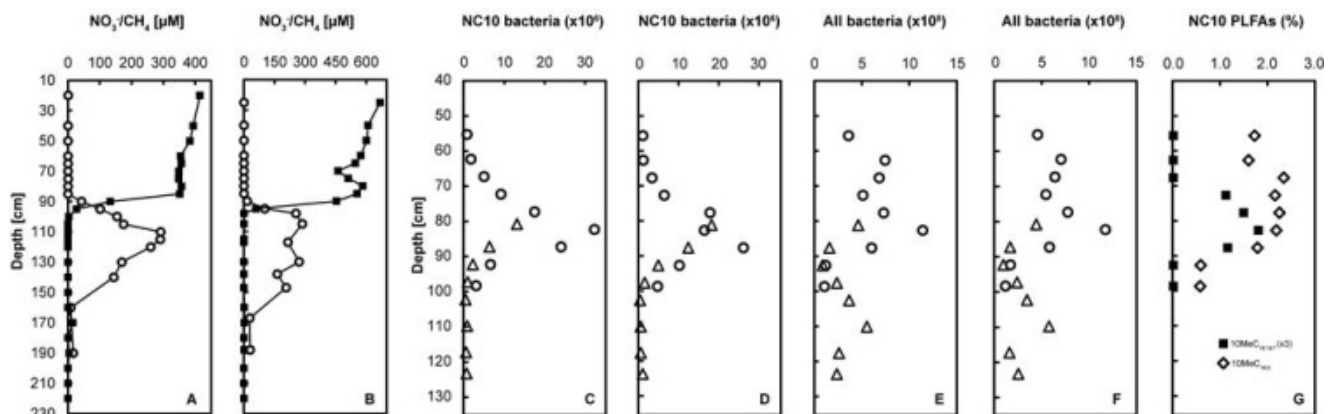
- 435 guidelines: minimum information for publication of quantitative real-time PCR experiments.  
436 Clin Chem 55:611-22.
- 437 5. Cicerone, R. J., and R. S. Oremland. 1988. Biogeochemical aspects of atmospheric methane.  
438 Global Biogeochem. Cycles 2:299-327.
- 439 6. Clymo, R. S., and C. L. Bryant. 2008. Diffusion and mass flow of dissolved carbon dioxide,  
440 methane, and dissolved organic carbon in a 7-m deep raised peat bog. Geochimica Et  
441 Cosmochimica Acta 72:2048-2066.
- 442 7. Daims, H., A. Bruhl, R. Amann, K. H. Schleifer, and M. Wagner. 1999. The domain-specific  
443 probe EUB338 is insufficient for the detection of all Bacteria: development and evaluation of  
444 a more comprehensive probe set. Syst Appl Microbiol 22:434-44.
- 445 8. Denman, K. L., G. Brasseur, A. Chidthaisong, P. Ciais, P. M. Cox, R. E. Dickinson, D.  
446 Hauglustaine, C. Heinze, E. Holland, D. Jacob, U. Lohmann, S. Ramachandran, P. L. da Silva  
447 Dias, S. C. Wofsy, and X. Zhang. 2007. Couplings Between Changes in the Climate System  
448 and Biogeochemistry. In S. Solomon, D. Qin, M. Manning, Z. Chen, M. Marquis, K. B.  
449 Averyt, M. Tignor, and H. L. Miller (ed.), Climate Change 2007: The Physical Science Basis.  
450 Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental  
451 Panel on Climate Change. Cambridge University Press.
- 452 9. Deutzmann, J. S., and B. Schink. 2011. Anaerobic oxidation of methane in sediments of Lake  
453 Constance, an oligotrophic freshwater lake. Appl Environ Microbiol 77:4429-36.
- 454 10. Downes, J., S. R. Vartoukian, F. E. Dewhirst, J. Izard, T. Chen, W. H. Yu, I. C. Sutcliffe, and  
455 W. G. Wade. 2009. *Pyramidobacter piscolens* gen. nov., sp nov., a member of the phylum  
456 'Synergistetes' isolated from the human oral cavity. International Journal of Systematic and  
457 Evolutionary Microbiology 59:972-980.
- 458 11. Eller, G., L. Kanel, and M. Krüger. 2005. Cooccurrence of aerobic and anaerobic methane  
459 oxidation in the water column of Lake Plusssee. Appl Environ Microbiol 71:8925-8.
- 460 12. Ettwig, K. F., M. K. Butler, D. Le Paslier, E. Pelletier, S. Mangenot, M. M. Kuypers, F.  
461 Schreiber, B. E. Dutilh, J. Zedelius, D. de Beer, J. Gloerich, H. J. Wessels, T. van Alen, F.  
462 Luesken, M. L. Wu, K. T. van de Pas-Schoonen, H. J. Op den Camp, E. M. Janssen-Megens,  
463 K. J. Francoijs, H. Stunnenberg, J. Weissenbach, M. S. Jetten, and M. Strous. 2010. Nitrite-  
464 driven anaerobic methane oxidation by oxygenic bacteria. Nature 464:543-8.
- 465 13. Ettwig, K. F., S. Shima, K. T. van de Pas-Schoonen, J. Kahnt, M. H. Medema, H. J. Op den  
466 Camp, M. S. Jetten, and M. Strous. 2008. Denitrifying bacteria anaerobically oxidize methane  
467 in the absence of Archaea. Environ Microbiol 10:3164-73.
- 468 14. Ettwig, K. F., T. van Alen, K. T. van de Pas-Schoonen, M. S. Jetten, and M. Strous. 2009.  
469 Enrichment and molecular detection of denitrifying methanotrophic bacteria of the NC10  
470 phylum. Appl Environ Microbiol 75:3656-62.
- 471 15. Fritz, C., V. A. Pancotto, J. T. Elzenga, E. J. Visser, A. P. Grootjans, A. Pol, R. Iturraspe, J. G.  
472 Roelofs, and A. J. Smolders. 2011. Zero methane emission bogs: extreme rhizosphere  
473 oxygenation by cushion plants in Patagonia. New Phytol 190:398-408.
- 474 16. Gärdenäs, A. I., G. I. Ågren, J. A. Bird, M. Clarholm, S. Hallin, P. Ineson, T. Kätterer, H.  
475 Knicker, S. I. Nilsson, T. Näsholm, S. Ogle, K. Paustian, T. Persson, and J. Stendahl. 2011.  
476 Knowledge gaps in soil carbon and nitrogen interactions – From molecular to global scale.  
477 Soil Biology and Biochemistry 43:702-717.

- 478 17. Gauci, V., N. Dise, and D. Fowler. 2002. Controls on suppression of methane flux from a peat  
479 bog subjected to simulated acid rain sulfate deposition. *Global Biogeochem. Cycles* 16:1004.
- 480 18. Gauci, V., E. Matthews, N. Dise, B. Walter, D. Koch, G. Granberg, and M. Vile. 2004. Sulfur  
481 pollution suppression of the wetland methane source in the 20th and 21st centuries. *Proc Natl*  
482 *Acad Sci U S A* 101:12583-7.
- 483 19. Hanson, R. S., and T. E. Hanson. 1996. Methanotrophic bacteria. *Microbiol Rev* 60:439-71.
- 484 20. Hu, S., R. J. Zeng, L. C. Burow, P. Lant, J. Keller, and Z. Yuan. 2009. Enrichment of  
485 denitrifying anaerobic methane oxidizing microorganisms. *Environmental Microbiology*  
486 *Reports* 1:377-384.
- 487 21. Juretschko, S., G. Timmermann, M. Schmid, K. H. Schleifer, A. Pommerening-Roser, H. P.  
488 Koops, and M. Wagner. 1998. Combined molecular and conventional analyses of nitrifying  
489 bacterium diversity in activated sludge: *Nitrosococcus mobilis* and *Nitrospira*-like bacteria as  
490 dominant populations. *Appl Environ Microbiol* 64:3042-51.
- 491 22. Kampman, C., T. L. G. Hendrickx, F. A. Luesken, T. A. van Alen, H. J. M. O. d. Camp, M. S.  
492 M. Jetten, G. Zeeman, C. J. N. Buisman, and H. Temmink. 2012. Enrichment of denitrifying  
493 methanotrophic bacteria for application after direct low-temperature anaerobic sewage  
494 treatment. *Journal of Hazardous Materials*.
- 495 23. Kartal, B., M. Koleva, R. Arsov, W. van der Star, M. S. Jetten, and M. Strous. 2006.  
496 Adaptation of a freshwater anammox population to high salinity wastewater. *J Biotechnol*  
497 126:546-53.
- 498 24. Kip, N., J. F. van Winden, Y. Pan, L. Bodrossy, G.-J. Reichart, A. J. P. Smolders, M. S. M.  
499 Jetten, J. S. Sinninghe Damsté, and H. J. M. Op den Camp. 2010. Global prevalence of  
500 methane oxidation by symbiotic bacteria in peat-moss ecosystems. *Nature Geosci* 3:617-621.
- 501 24A. Kool DM, Zhu B, et al. 2012. Rare branched fatty acids characterize the lipid composition of  
502 the intra-aerobic methane oxidizer *Candidatus* *Mehtylomirabilis oxyfera*. *Appl. Environ.*  
503 *Microbiol.* 78:8650-8656.
- 504 25. Knittel, K., and A. Boetius. 2009. Anaerobic oxidation of methane: progress with an unknown  
505 process. *Annu Rev Microbiol* 63:311-34.
- 506 26. Lai, D. Y. F. 2009. Methane Dynamics in Northern Peatlands: A Review. *Pedosphere* 19:409-  
507 421.
- 508 27. Le Mer, J., and P. Roger. 2001. Production, oxidation, emission and consumption of methane  
509 by soils: A review. *European Journal of Soil Biology* 37:25-50.
- 510 28. Leser, T. D., J. Z. Amenuvor, T. K. Jensen, R. H. Lindecrone, M. Boye, and K. Moller. 2002.  
511 Culture-independent analysis of gut bacteria: the pig gastrointestinal tract microbiota revisited.  
512 *Appl Environ Microbiol* 68:673-90.
- 513 29. Liu, L. L., and T. L. Greaver. 2009. A review of nitrogen enrichment effects on three biogenic  
514 GHGs: the CO<sub>2</sub> sink may be largely offset by stimulated N<sub>2</sub>O and CH<sub>4</sub> emission. *Ecology*  
515 *Letters* 12:1103-1117.
- 516 30. Lovley, D. R., and M. J. Klug. 1983. Sulfate Reducers Can Outcompete Methanogens at  
517 Freshwater Sulfate Concentrations. *Applied and Environmental Microbiology* 45:187-192.
- 518 31. Lovley, D. R., and E. J. Phillips. 1987. Competitive mechanisms for inhibition of sulfate  
519 reduction and methane production in the zone of ferric iron reduction in sediments. *Appl*  
520 *Environ Microbiol* 53:2636-41.

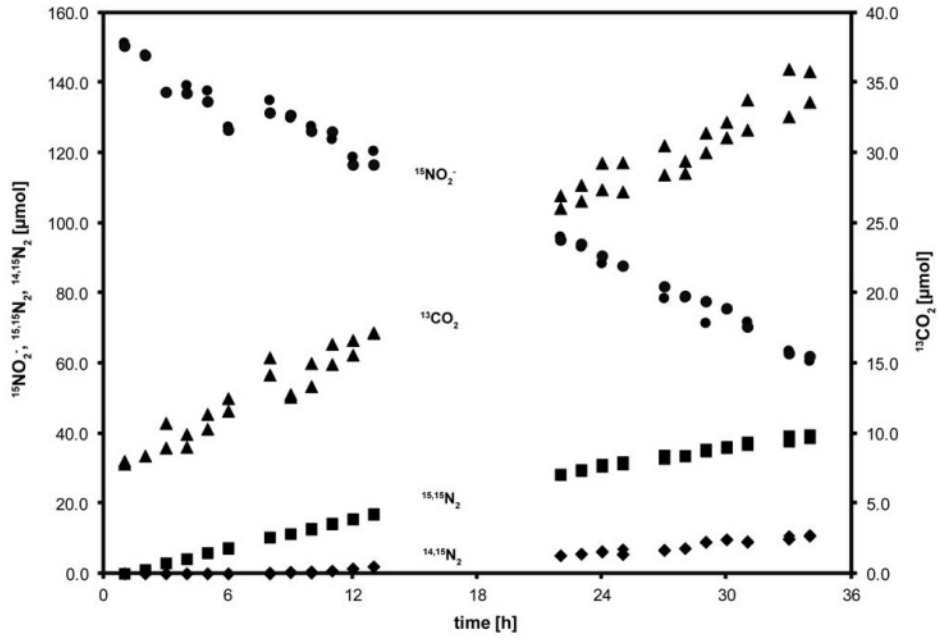
- 521 32. Lucassen, E. C. H. E. T., A. J. P. Smolders, A. L. Van der Salm, and J. G. M. Roelofs. 2004.  
522 High groundwater nitrate concentrations inhibit eutrophication of sulphate-rich freshwater  
523 wetlands. *Biogeochemistry* 67:249-267.
- 524 33. Luesken, F. A., T. A. van Alen, E. van der Biezen, C. Frijters, G. Toonen, C. Kampman, T. L.  
525 Hendrickx, G. Zeeman, H. Temmink, M. Strous, H. J. Op den Camp, and M. S. Jetten. 2011.  
526 Diversity and enrichment of nitrite-dependent anaerobic methane oxidizing bacteria from  
527 wastewater sludge. *Appl Microbiol Biotechnol*.
- 528 34. Luesken, F. A., B. Zhu, T. A. van Alen, M. K. Butler, M. R. Diaz, B. Song, H. J. Op den  
529 Camp, M. S. Jetten, and K. F. Ettwig. 2011. *pmoA* Primers for detection of anaerobic  
530 methanotrophs. *Appl Environ Microbiol* 77:3877-80.
- 531 35. Lüke, C., and P. Frenzel. 2011. Potential of *pmoA* Amplicon Pyrosequencing for  
532 Methanotroph Diversity Studies. *Applied and Environmental Microbiology* 77:6305-6309.
- 533 36. McDonald, I. R., L. Bodrossy, Y. Chen, and J. C. Murrell. 2008. Molecular ecology  
534 techniques for the study of aerobic methanotrophs. *Appl Environ Microbiol* 74:1305-15.
- 535 37. Pester, M., K. H. Knorr, M. W. Friedrich, M. Wagner, and A. Loy. 2012. Sulfate-reducing  
536 microorganisms in wetlands - fameless actors in carbon cycling and climate change. *Front*  
537 *Microbiol* 3:72.
- 538 38. Raghoebarsing, A. A., A. Pol, K. T. van de Pas-Schoonen, A. J. Smolders, K. F. Ettwig, W. I.  
539 Rijpstra, S. Schouten, J. S. Sinninghe Damsté, H. J. Op den Camp, M. S. Jetten, and M.  
540 Strous. 2006. A microbial consortium couples anaerobic methane oxidation to denitrification.  
541 *Nature* 440:918-21.
- 542 39. Reeburgh, W. S. 2007. Oceanic methane biogeochemistry. *Chem Rev* 107:486-513.
- 543 40. Schubert, C. J., F. Vazquez, T. Lösekann-Behrens, K. Knittel, M. Tonolla, and A. Boetius.  
544 2011. Evidence for anaerobic oxidation of methane in sediments of a freshwater system (Lago  
545 di Cadagno). *FEMS Microbiol Ecol* 76:26-38.
- 546 41. Segers, R. 1998. Methane production and methane consumption: a review of processes  
547 underlying wetland methane fluxes. *Biogeochemistry* 41:23-51.
- 548 42. Sivan, O., M. Adler, A. Pearson, F. Gelman, I. Bar-Or, S. G. John, and W. Eckert. 2011.  
549 Geochemical evidence for iron-mediated anaerobic oxidation of methane. *Limnology and*  
550 *Oceanography* 56:1536-1544.
- 551 43. Smemo, K. A., and J. B. Yavitt. 2011. Anaerobic oxidation of methane: an underappreciated  
552 aspect of methane cycling in peatland ecosystems? *Biogeosciences* 8:779-793.
- 553 44. Smith, R. L., B. L. Howes, and S. P. Garabedian. 1991. In situ measurement of methane  
554 oxidation in groundwater by using natural-gradient tracer tests. *Appl Environ Microbiol*  
555 57:1997-2004.
- 556 45. Smolders, A. J. P., E. C. H. E. T. Lucassen, R. Bobbink, J. G. M. Roelofs, and L. P. M.  
557 Lamers. 2010. How nitrate leaching from agricultural lands provokes phosphate  
558 eutrophication in groundwater fed wetlands: the sulphur bridge. *Biogeochemistry* 98:1-7.
- 559 46. Stevens, C. J., P. Manning, L. J. van den Berg, M. C. de Graaf, G. W. Wamelink, A. W.  
560 Boxman, A. Bleeker, P. Vergeer, M. Arroniz-Crespo, J. Limpens, L. P. Lamers, R. Bobbink,  
561 and E. Dorland. 2011. Ecosystem responses to reduced and oxidised nitrogen inputs in  
562 European terrestrial habitats. *Environ Pollut* 159:665-76.



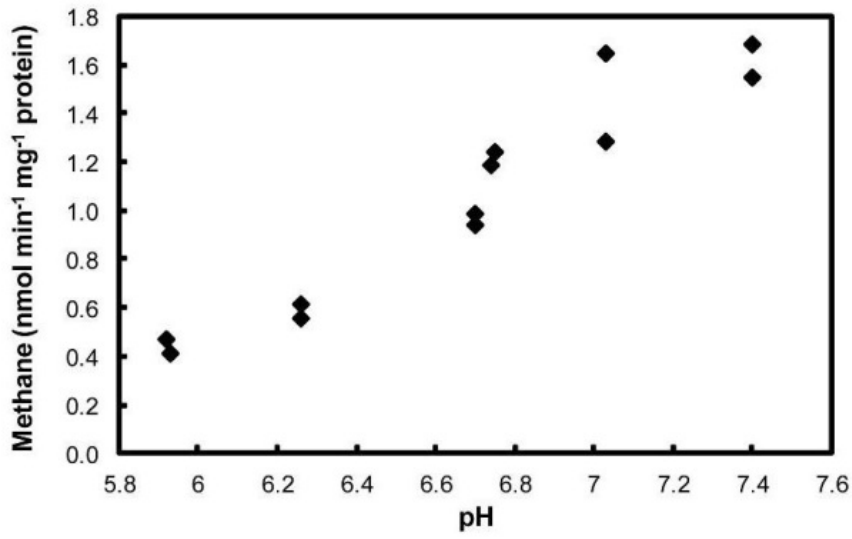
- 563 47. Strous, M., and M. S. Jetten. 2004. Anaerobic oxidation of methane and ammonium. *Annu*  
564 *Rev Microbiol* 58:99-117.
- 565 48. Tomassen, H. B. M., A. J. P. Smolders, P. M. L. Leon, and J. G. M. Roelofs. 2003. Stimulated  
566 Growth of *Betula pubescens* and *Molinia caerulea* on Ombrotrophic Bogs: Role of High  
567 Levels of Atmospheric Nitrogen Deposition. *Journal of Ecology* 91:357-370.
- 568 49. Turner, S., K. M. Pryer, V. P. W. Miao, and J. D. Palmer. 1999. Investigating deep  
569 phylogenetic relationships among cyanobacteria and plastids by small submit rRNA sequence  
570 analysis. *Journal of Eukaryotic Microbiology* 46:327-338.
- 571 50. van Breukelen, B. M., and J. Griffioen. 2004. Biogeochemical processes at the fringe of a  
572 landfill leachate pollution plume: potential for dissolved organic carbon, Fe(II), Mn(II), NH<sub>4</sub>,  
573 and CH<sub>4</sub> oxidation. *Journal of Contaminant Hydrology* 73:181-205.
- 574 51. Vile, M. A., S. D. Bridgham, R. K. Wieder, and M. Novák. 2003. Atmospheric sulfur  
575 deposition alters pathways of gaseous carbon production in peatlands. *Global Biogeochem.*  
576 *Cycles* 17:1058.
- 577 52. Waughman, G. J. 1980. Chemical Aspects of the Ecology of Some South German Peatlands.  
578 *Journal of Ecology* 68:1025-1046.
- 579 53. Weisburg, W. G., S. M. Barns, D. A. Pelletier, and D. J. Lane. 1991. 16s Ribosomal DNA  
580 Amplification for Phylogenetic Study. *Journal of Bacteriology* 173:697-703.
- 581 54. Zhu, G., M. S. Jetten, P. Kusch, K. F. Ettwig, and C. Yin. 2010. Potential roles of anaerobic  
582 ammonium and methane oxidation in the nitrogen cycle of wetland ecosystems. *Appl*  
583 *Microbiol Biotechnol* 86:1043-55.
- 584
- 585
- 586



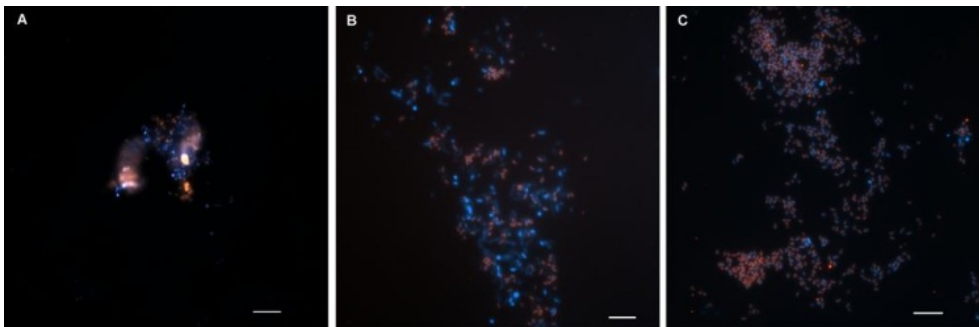
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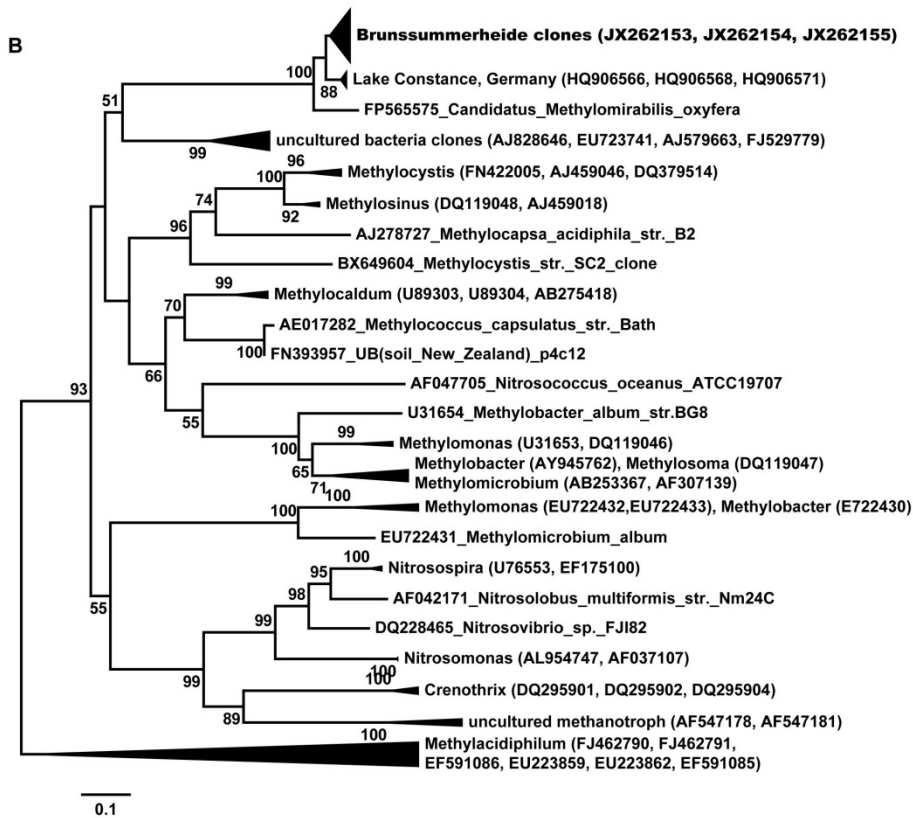
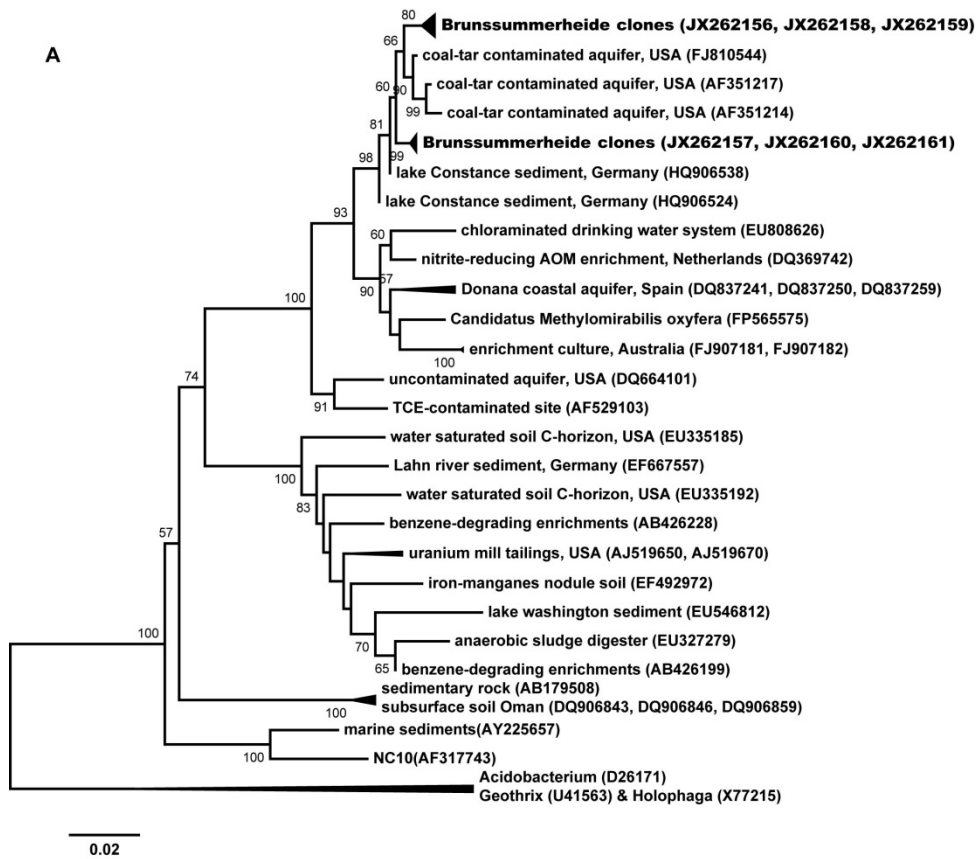
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599 Fig. 5